

HEAVY RAINFALL EVENTS PRECEDING THE ARRIVAL OF TROPICAL CYCLONES

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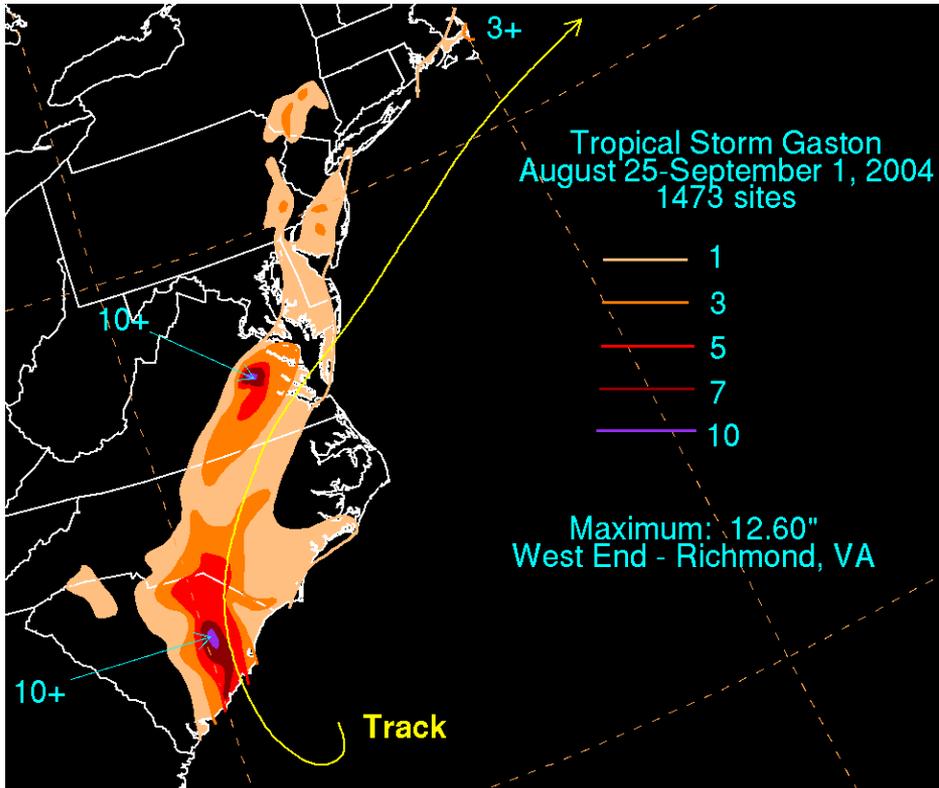
*National Weather Service Forecast Office
Binghamton, NY*

CSTAR II Grant NA04NWS4680005

Spring CSTAR II Meeting

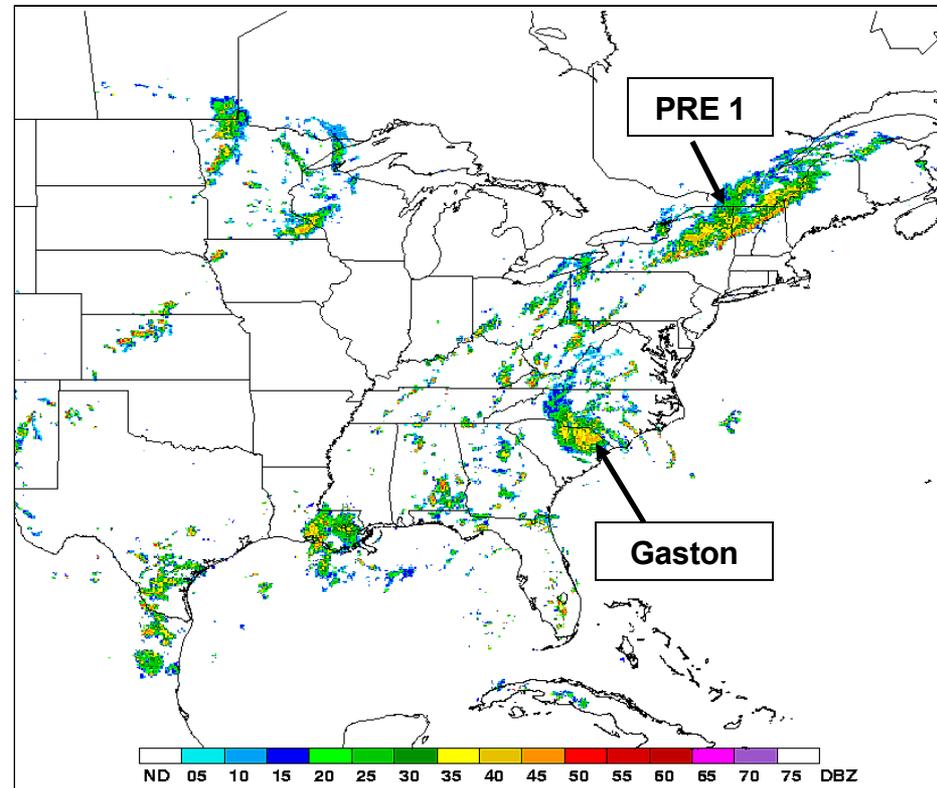
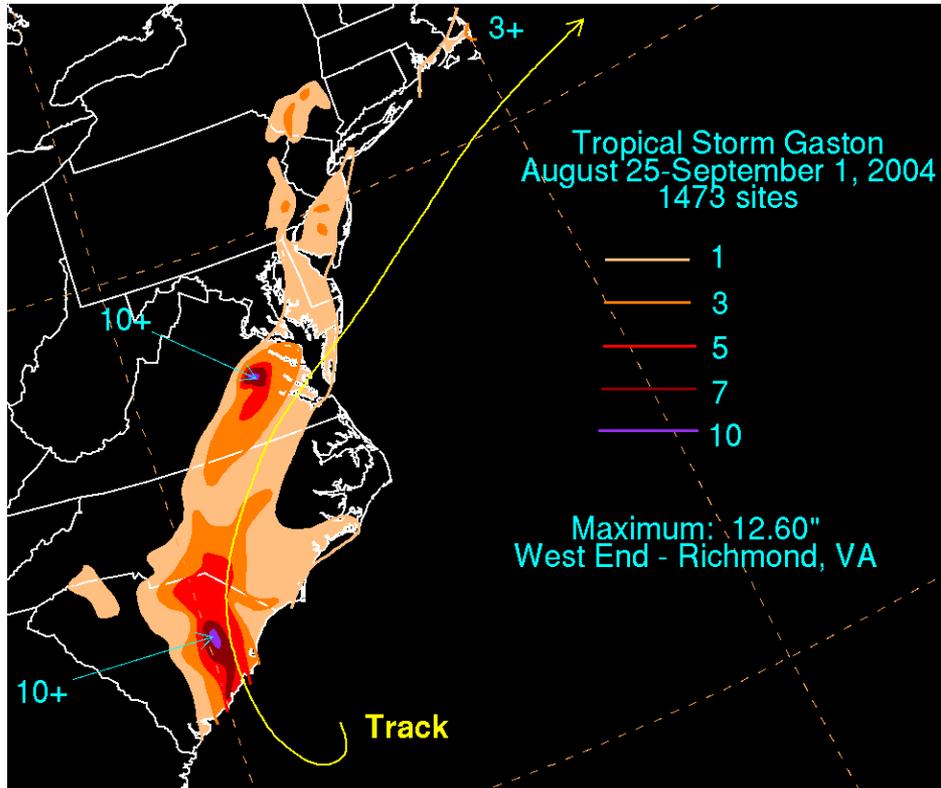
4 May 2007

INTRODUCTION



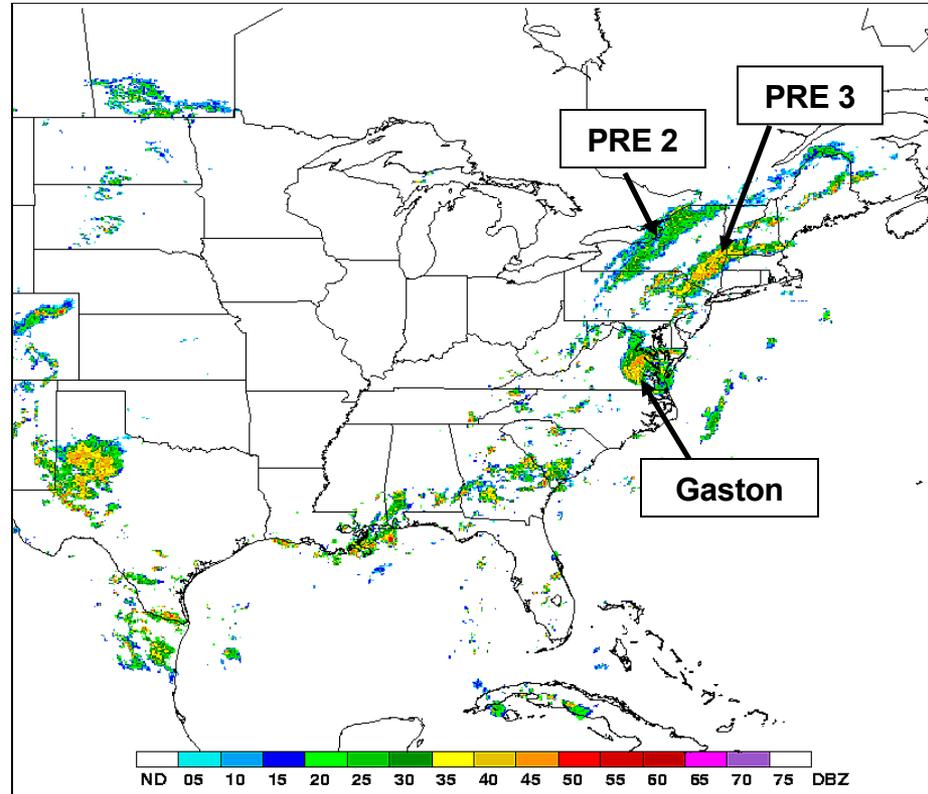
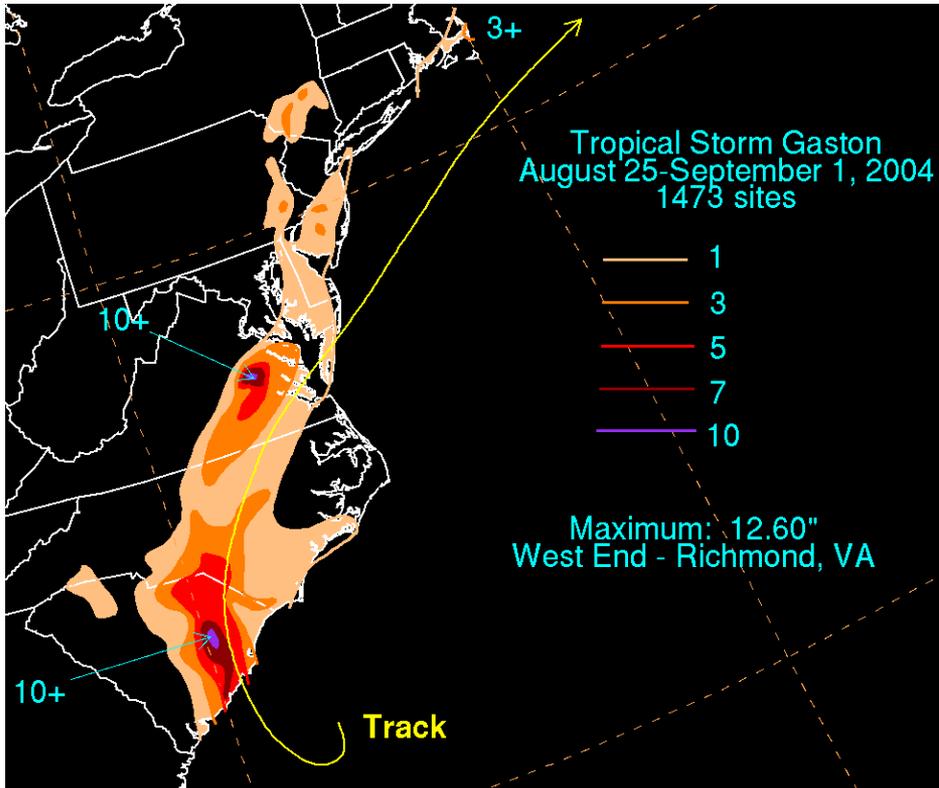
- Nothing unusual about rainfall distribution at first glance
- Heaviest in SC and VA just left of the TC track
- A few scattered areas of heavy rain further north

INTRODUCTION



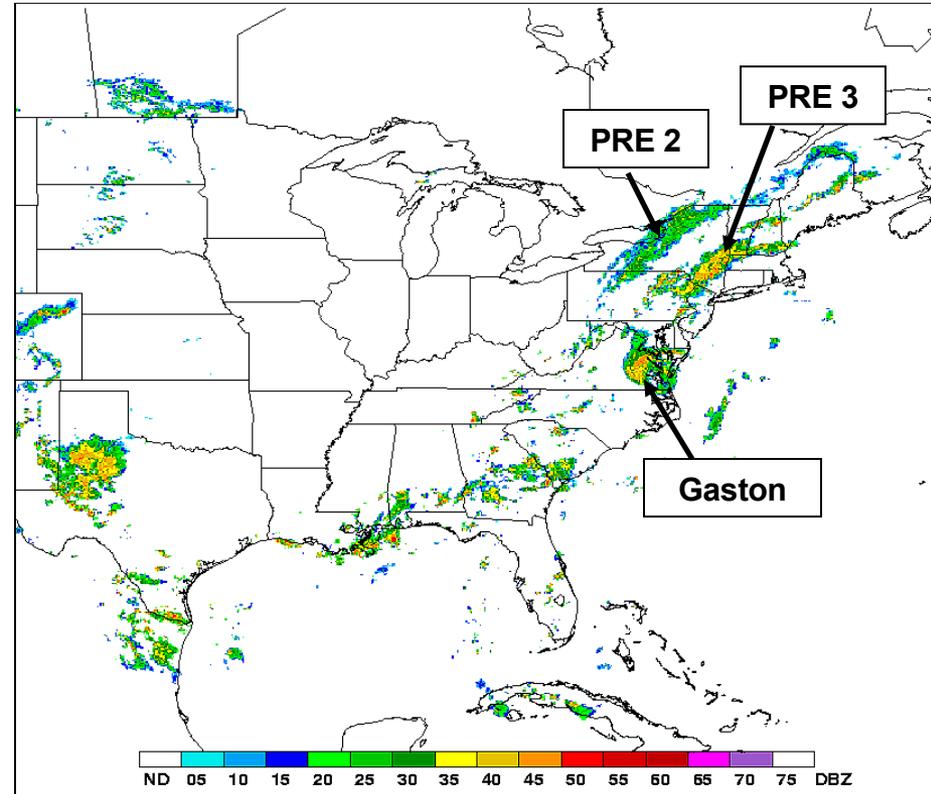
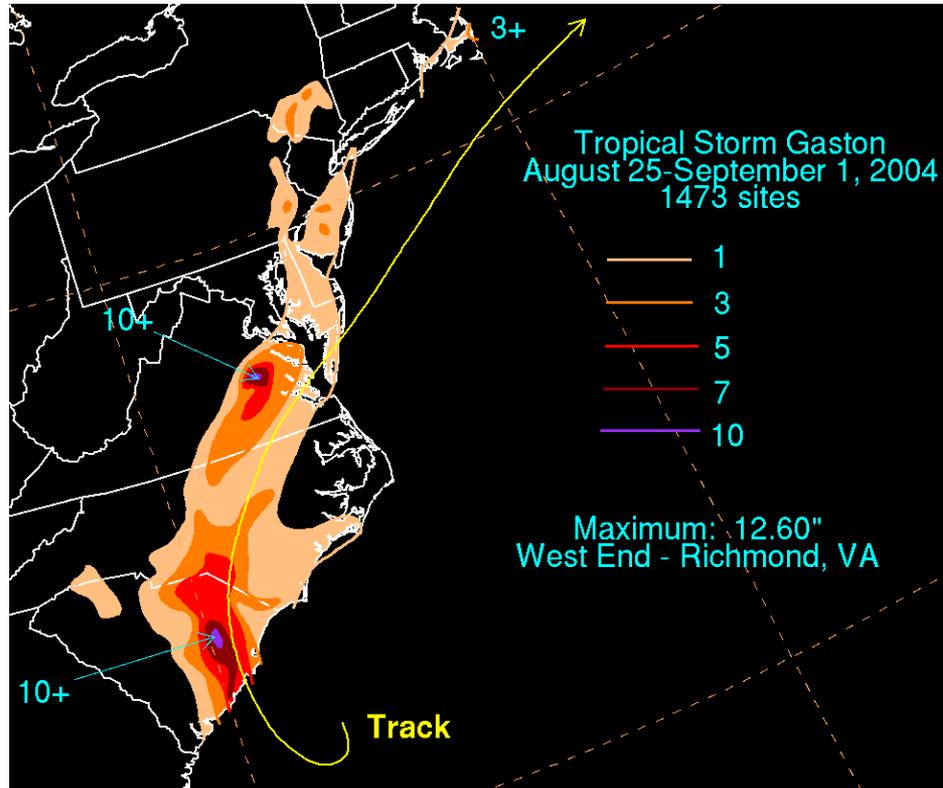
0000 UTC 040830 WSI NOWRAD Radar Mosaic

INTRODUCTION



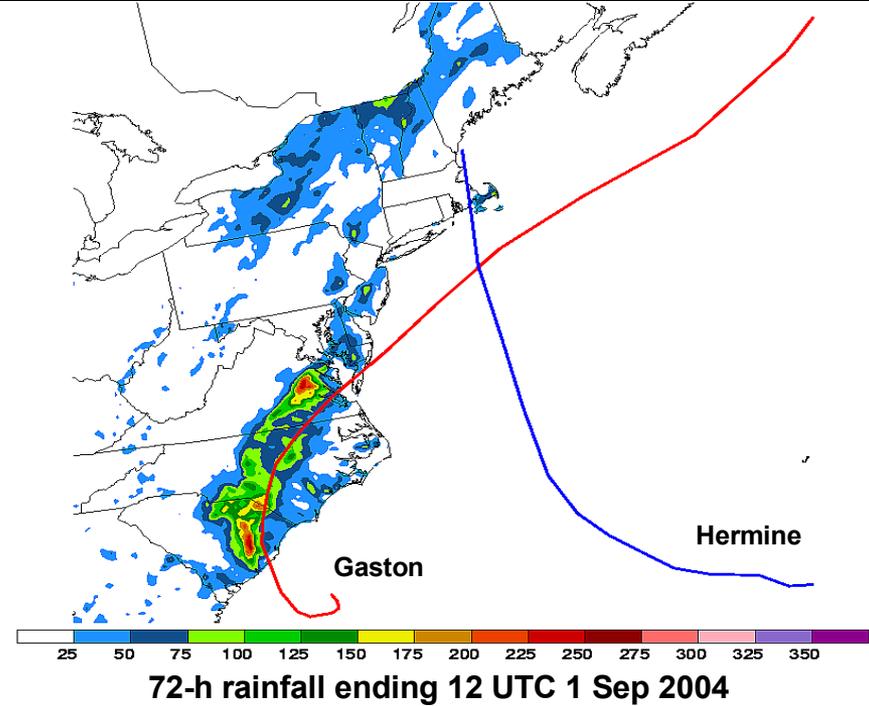
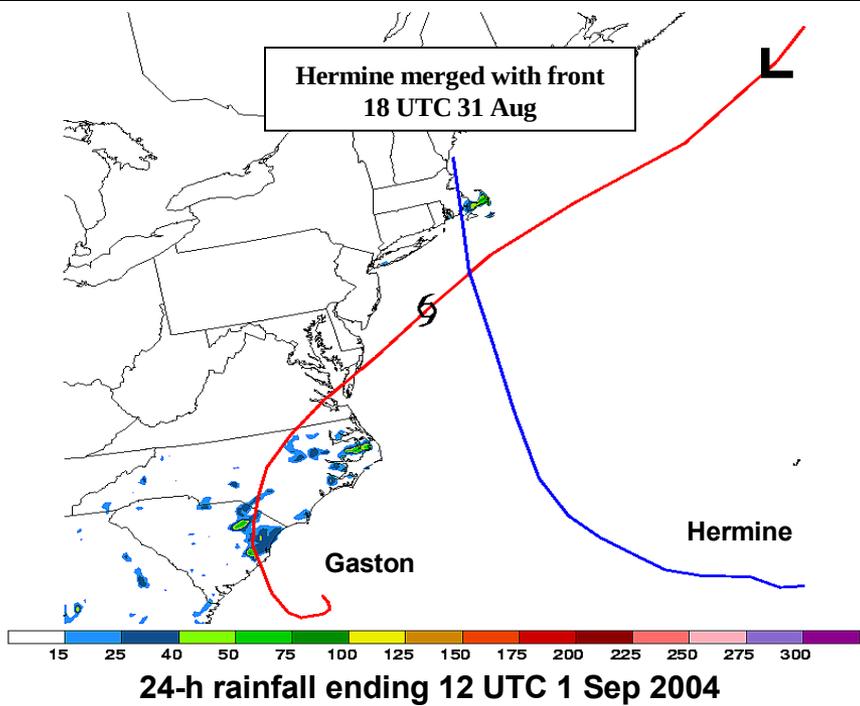
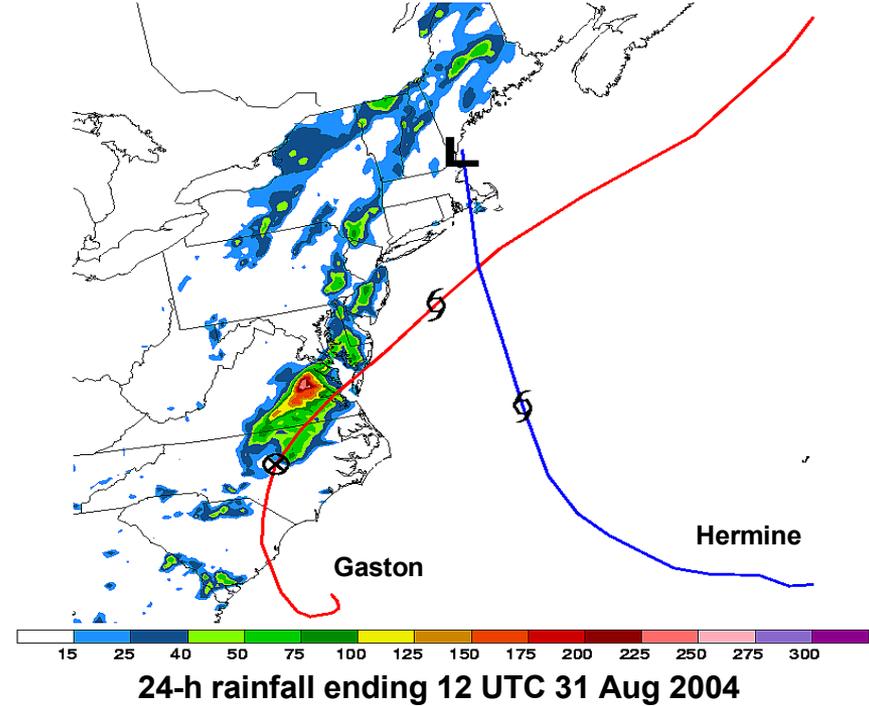
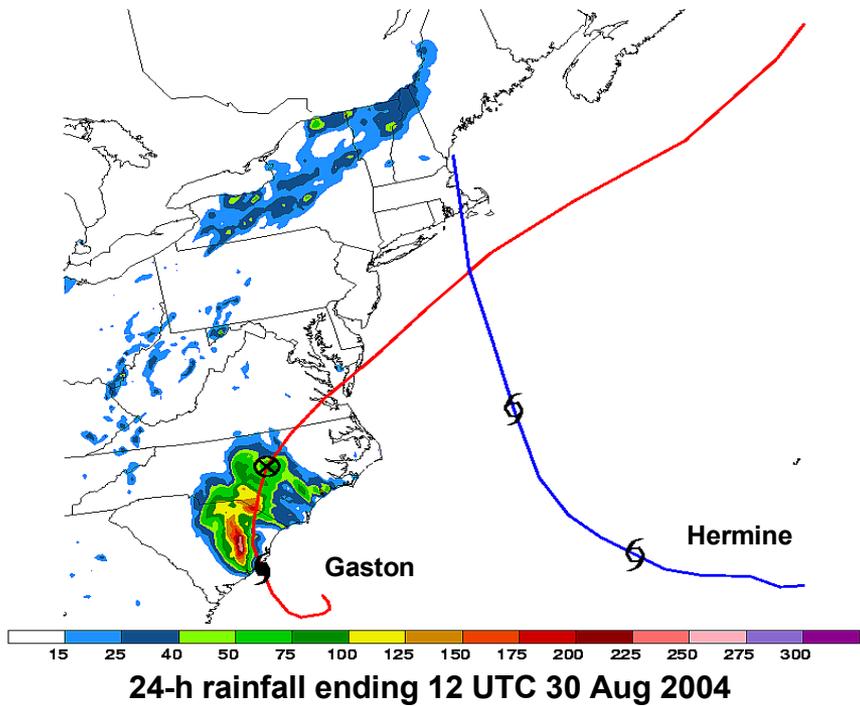
0000 UTC 040831 WSI NOWRAD Radar Mosaic

INTRODUCTION



0000 UTC 040831 WSI NOWRAD Radar Mosaic

How can we capture the downstream rainfall?



DATA SOURCES

- **NCDC and WSI NOWRAD radar imagery**
- **NHC best-track data**
- **NPVU QPE and NWS rainfall products**
- **NCEP/NARR gridded datasets**
- **NCEP/NCAR Global Reanalysis for compositing**
- **DATSAV and COADS surface data**

IDENTIFYING PREs (1998–2006)

- **Coherent area of rain displaced poleward of TC**
- **Maximum rainfall rates exceeded 100 mm in 24**
- **Moisture transport from TC toward PRE**

IDENTIFYING PREs (1998–2006)

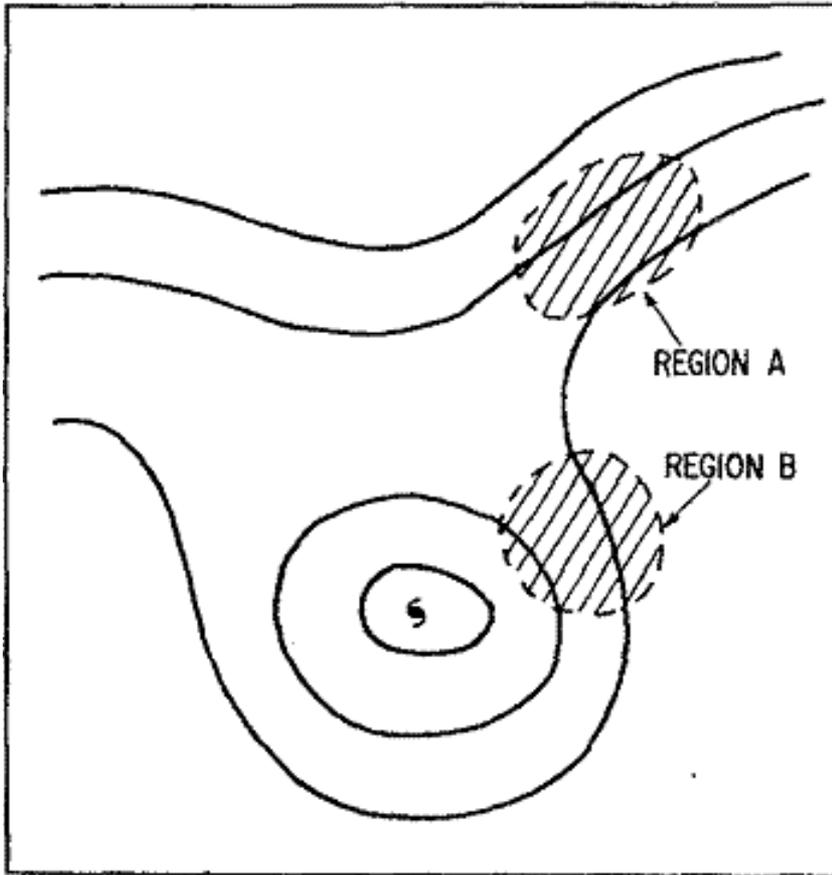
- Coherent area of rain displaced poleward of TC
- Maximum rainfall rates exceeded 100 mm in 24 h
- Moisture transport from TC toward PRE

**47 PREs associated with 21 TCs were identified
(~ 2 PREs per TC)**

**~ 1/3 of all US landfalling TCs produced at least one
PRE**

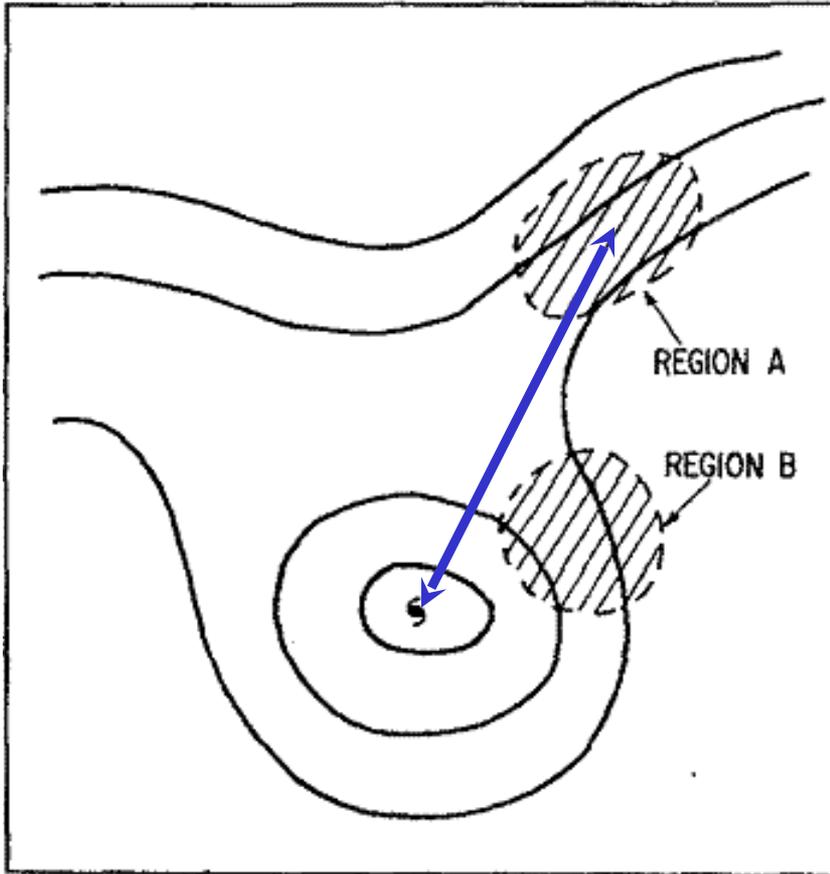
5 cases where TC did not make US landfall

PAST RESEARCH ON PRE WITH AGNES (1972)



Bosart and Carr (1978) conceptual model of antecedent rainfall

PRE STATISTICS



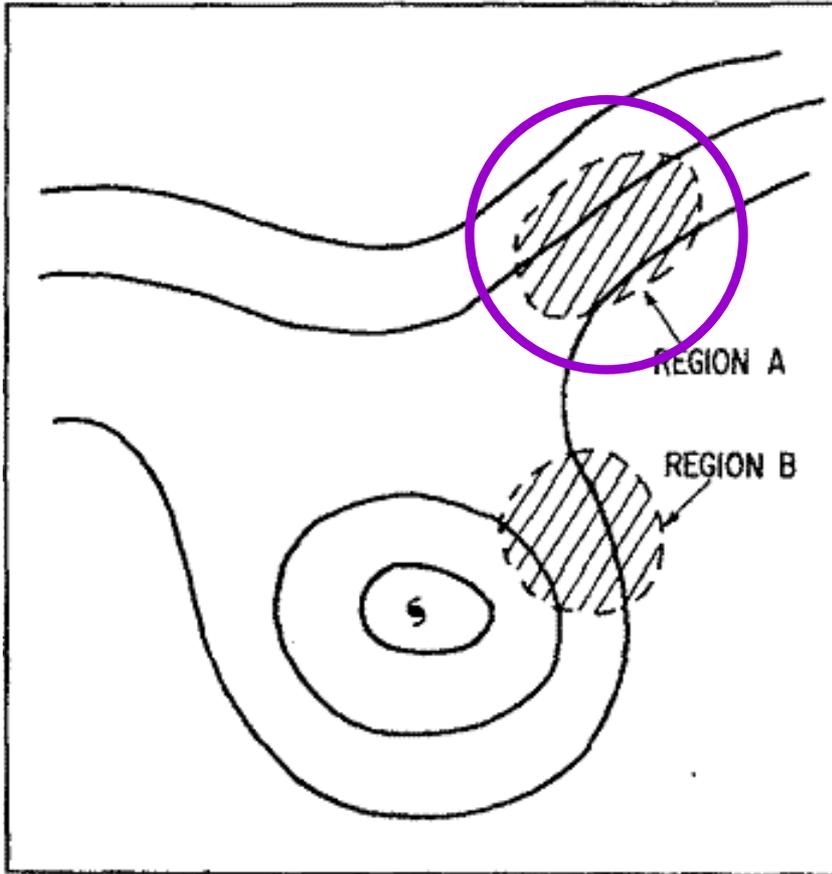
Separation Distance

1086 ± 482 km

Median: 935 km

**Bosart and Carr (1978) conceptual
model of antecedent rainfall**

PRE STATISTICS



Bosart and Carr (1978) conceptual model of antecedent rainfall

Separation Distance

1086 ± 482 km

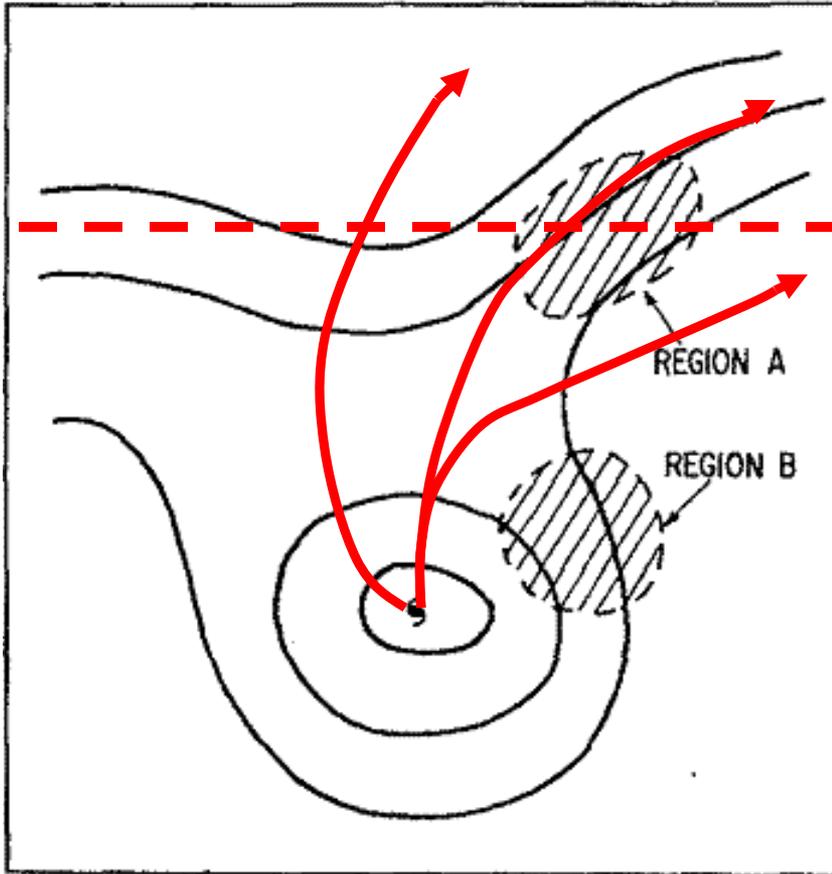
Median: 935 km

Event Duration

14 ± 7 h

Median: 12 h

PRE STATISTICS



Bosart and Carr (1978) conceptual model of antecedent rainfall

Separation Distance

1086 ± 482 km

Median: 935 km

Event Duration

14 ± 7 h

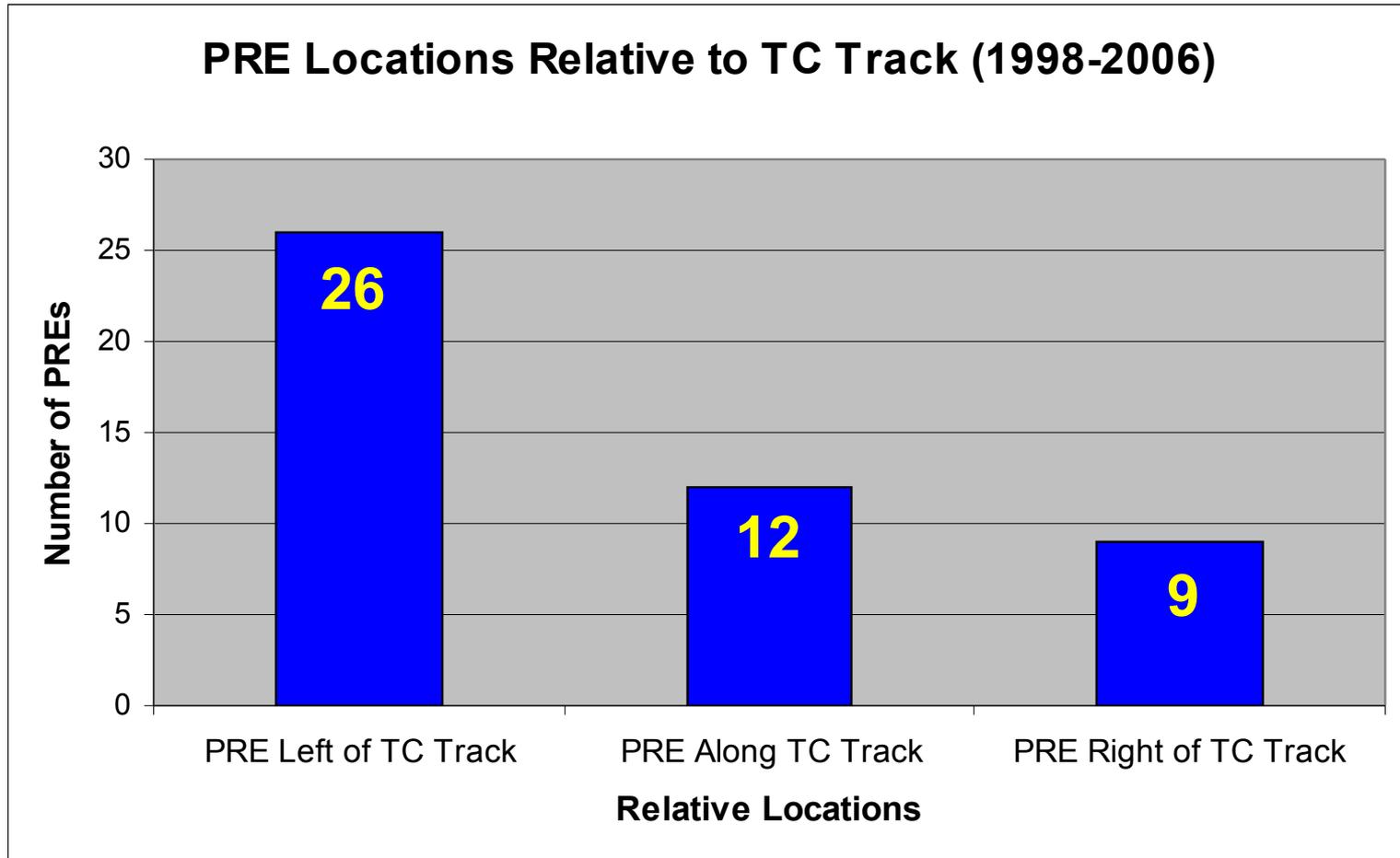
Median: 12 h

Time Lag

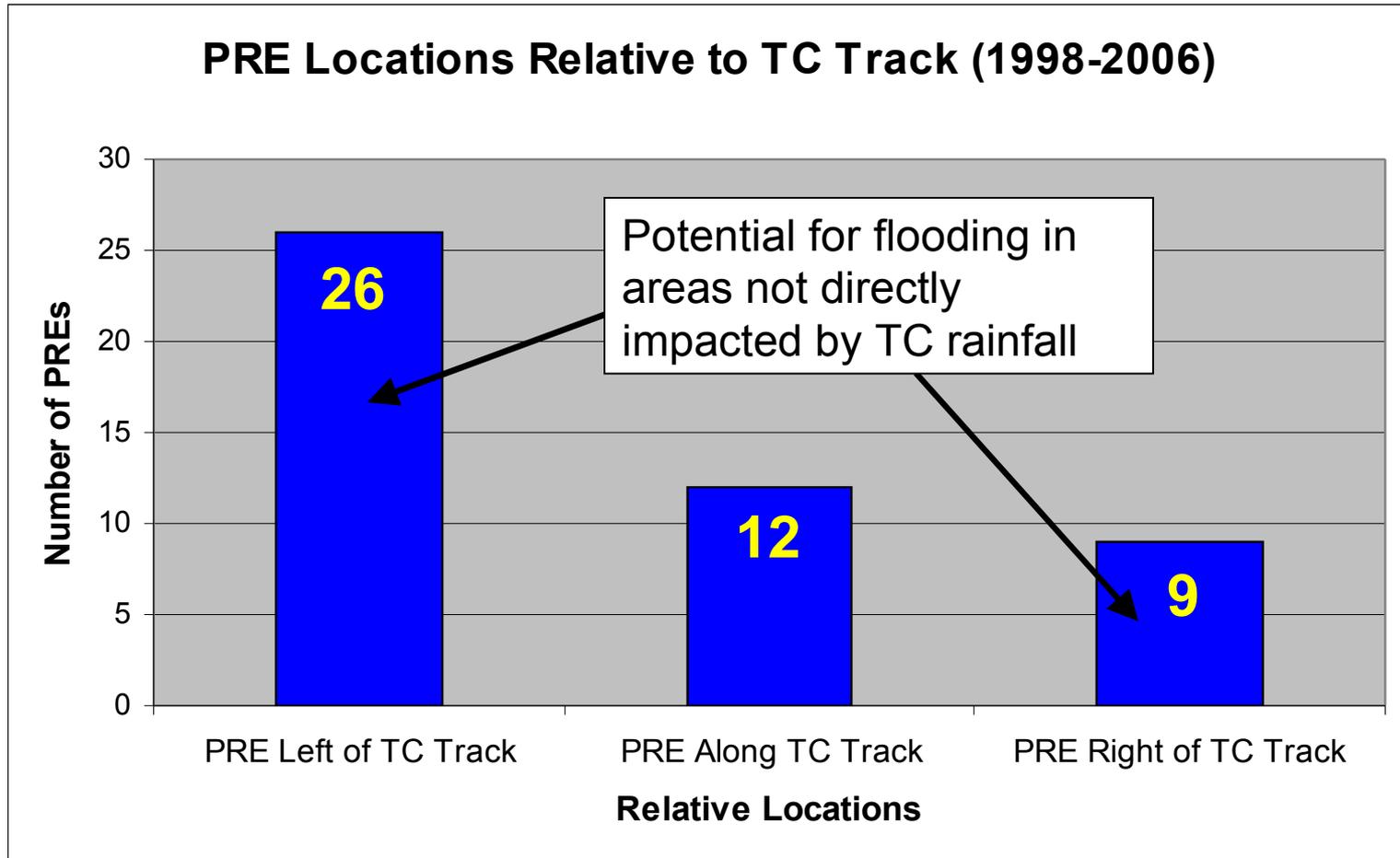
45 ± 29 h

Median: 36 h

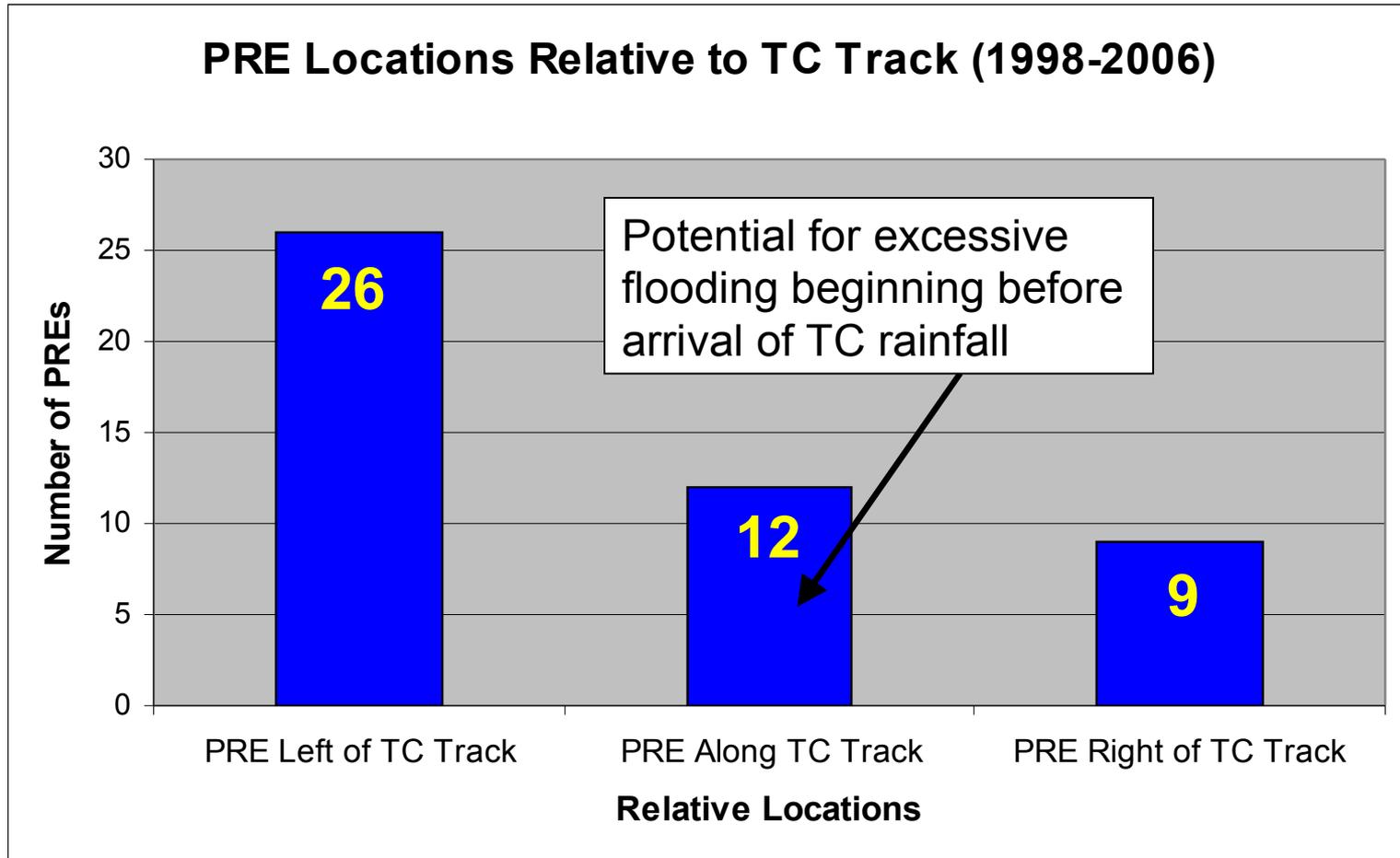
PRE TRACK-RELATIVE POSITIONS



PRE TRACK-RELATIVE POSITIONS



PRE TRACK-RELATIVE POSITIONS



PRE TRACK-RELATIVE POSITIONS

Type of PRE (Number in category)	24-h rainfall rate statistics (mm)			Mean PRE speed (m s ⁻¹)
	Mean	Std. deviation	Maximum	
Left of Track (22)	185	70	340	10.7
Along Track (8)	245	100	410	12.9
Right of Track (7)	260	80	410	5.7

GREATEST RAINFALL

SLOWEST MOVEMENT

PRE TRACK-RELATIVE POSITIONS

Type of PRE (Number in category)	24-h rainfall rate statistics (mm)			Mean PRE speed (m s ⁻¹)
	Mean	Std. deviation	Maximum	
Left of Track (22)	185	70	340	10.7
Along Track (8)	245	100	410	12.9
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HIGH RAINFALL

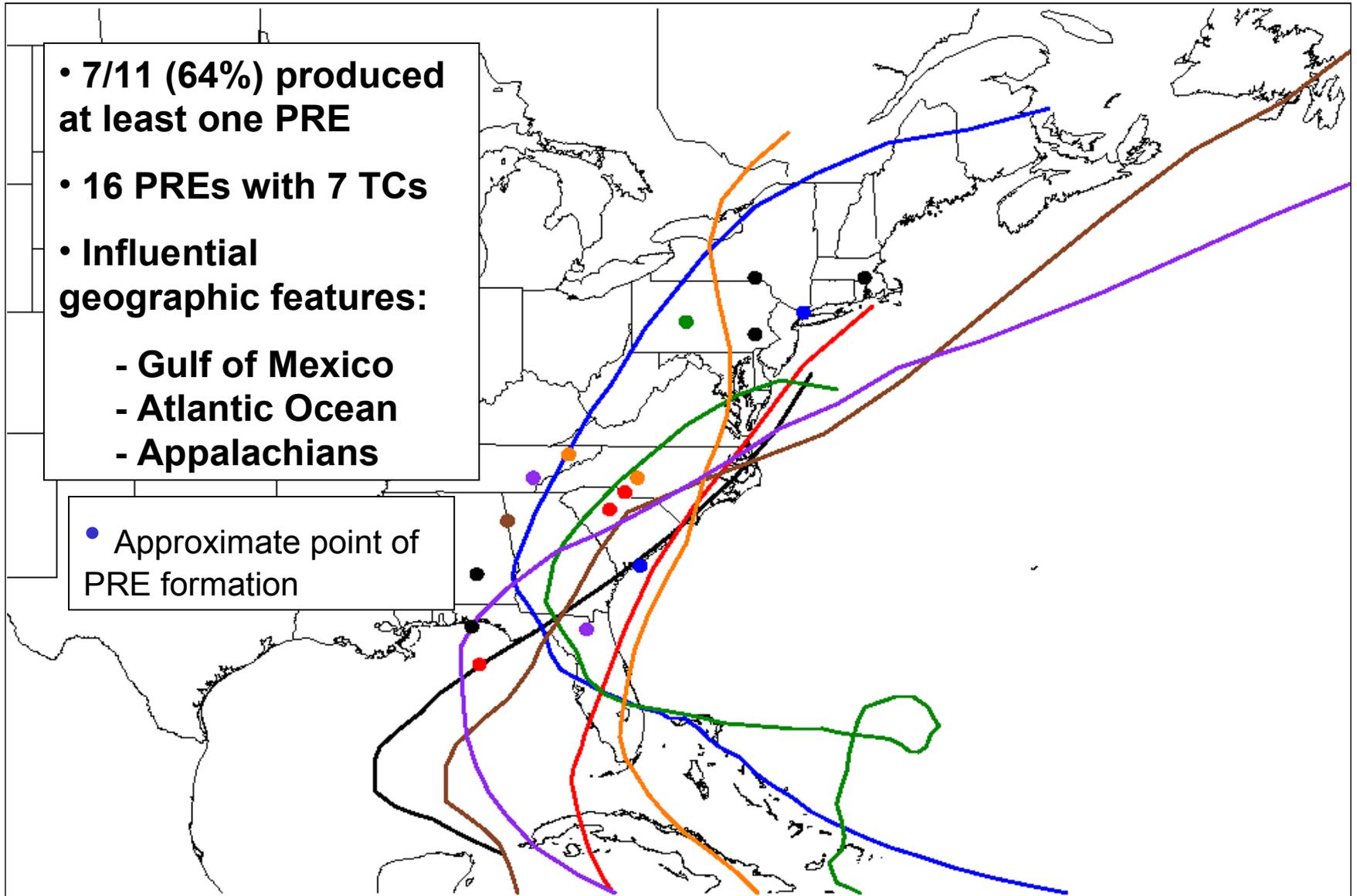
PREs MOVE TWICE AS FAST

SEPARATION BY TC TRACK SIMILARITY

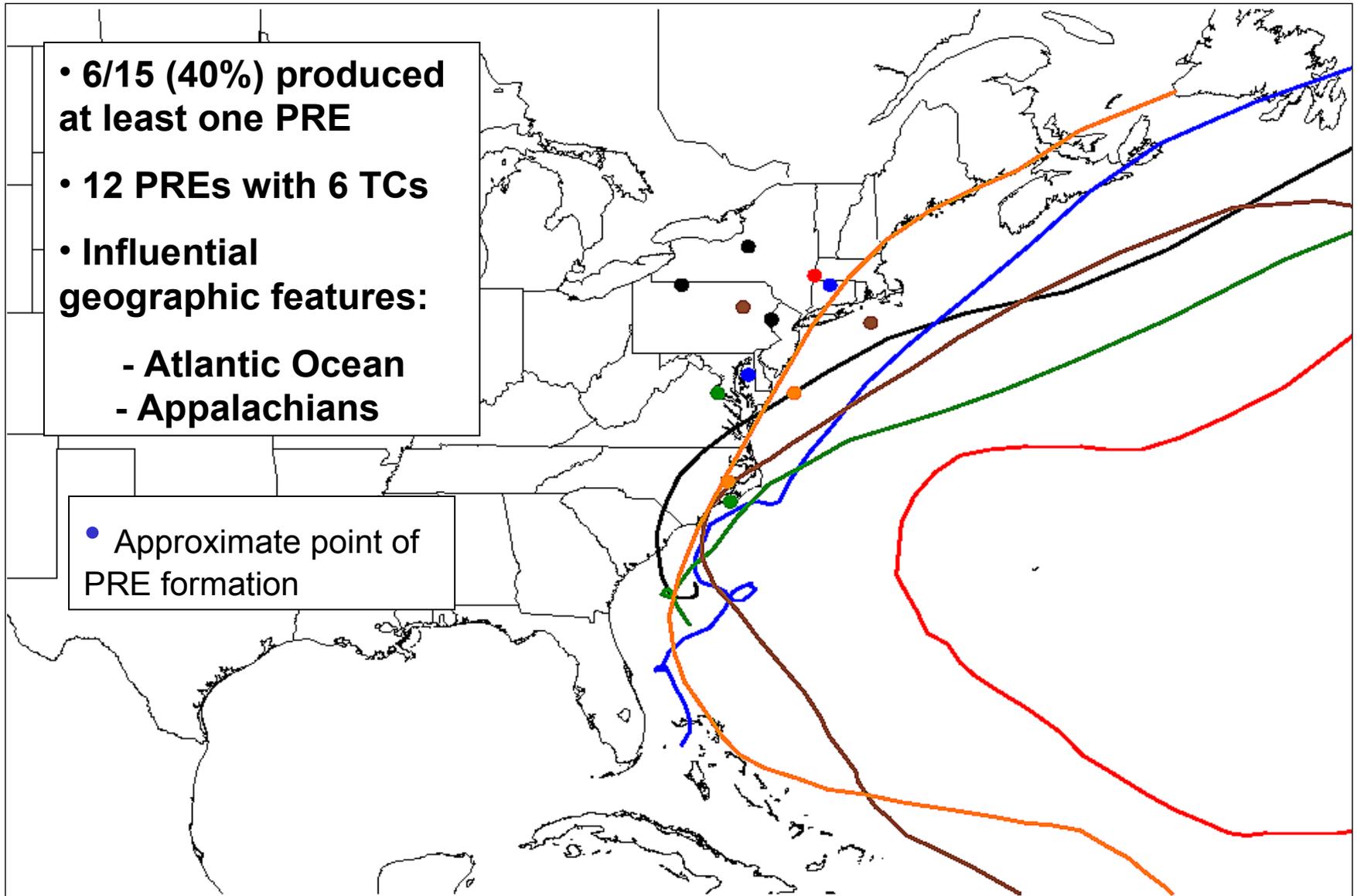
SOUTHEAST RECURVATURES

- 7/11 (64%) produced at least one PRE
- 16 PREs with 7 TCs
- Influential geographic features:
 - Gulf of Mexico
 - Atlantic Ocean
 - Appalachians

• Approximate point of PRE formation

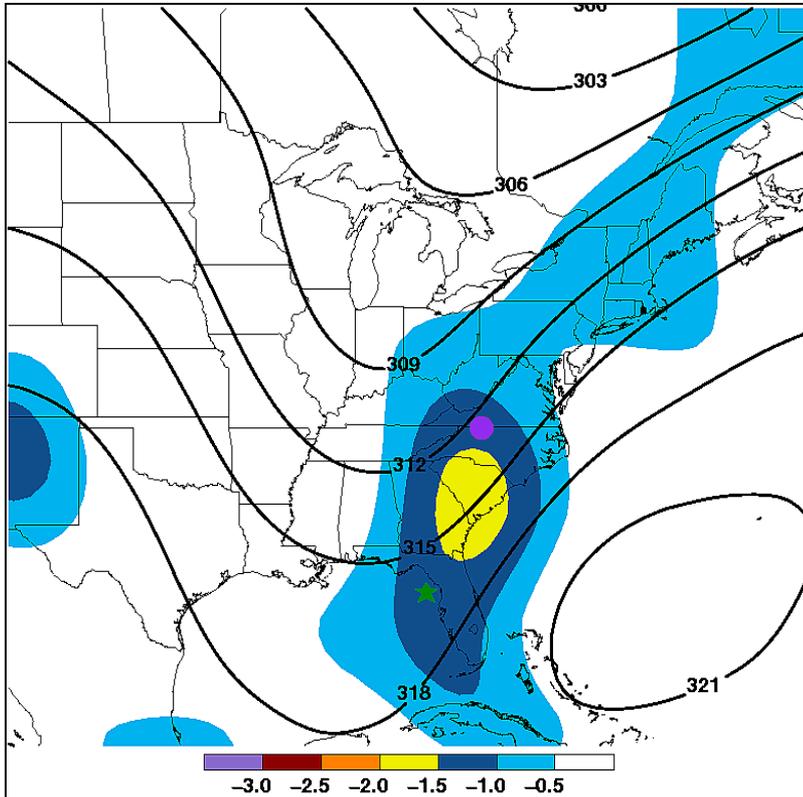


ATLANTIC RECURVATURES



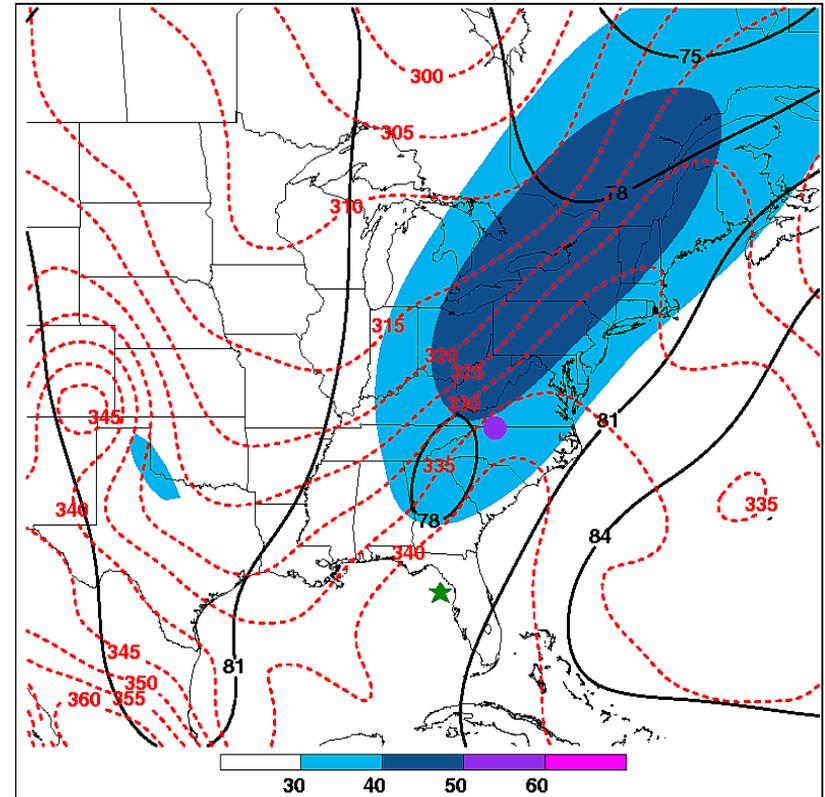
SE RECURVATURE PRE COMPOSITE

TIME OF PRE INITIATION



700 hPa Ht (dam) and UVM ($\mu\text{b s}^{-1}$)

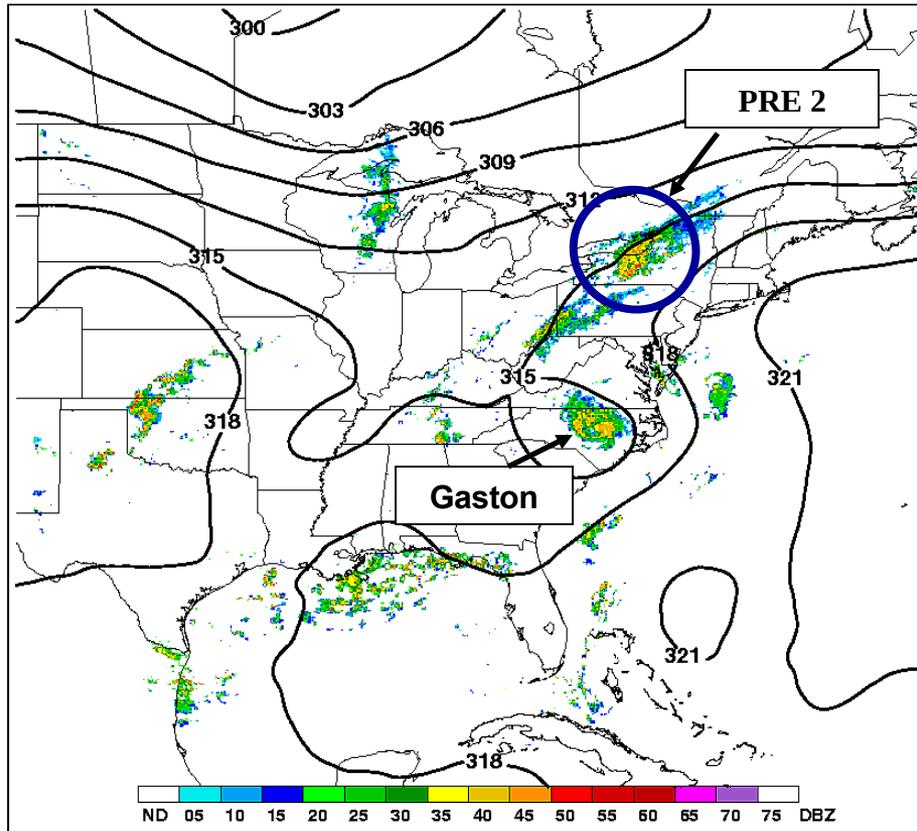
- Significant midlevel trough with weak UVM well poleward of TC
- Deep meridional flow transports tropical moisture up East Coast



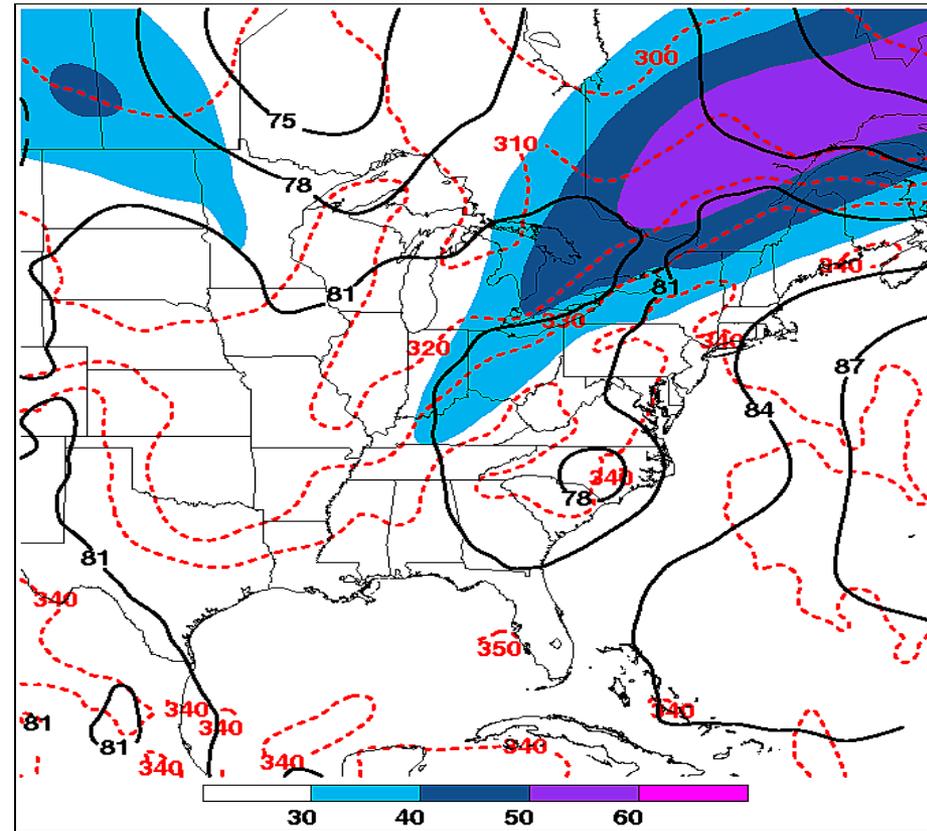
925 hPa Ht (dam), θ_e (K), and 200 hPa wind speeds (m s^{-1})

- PRE forms:
 - in right-entrance region of intensifying upper-level jet
 - on western edge of θ_e ridge

GASTON (2004): SYNOPTIC FEATURES

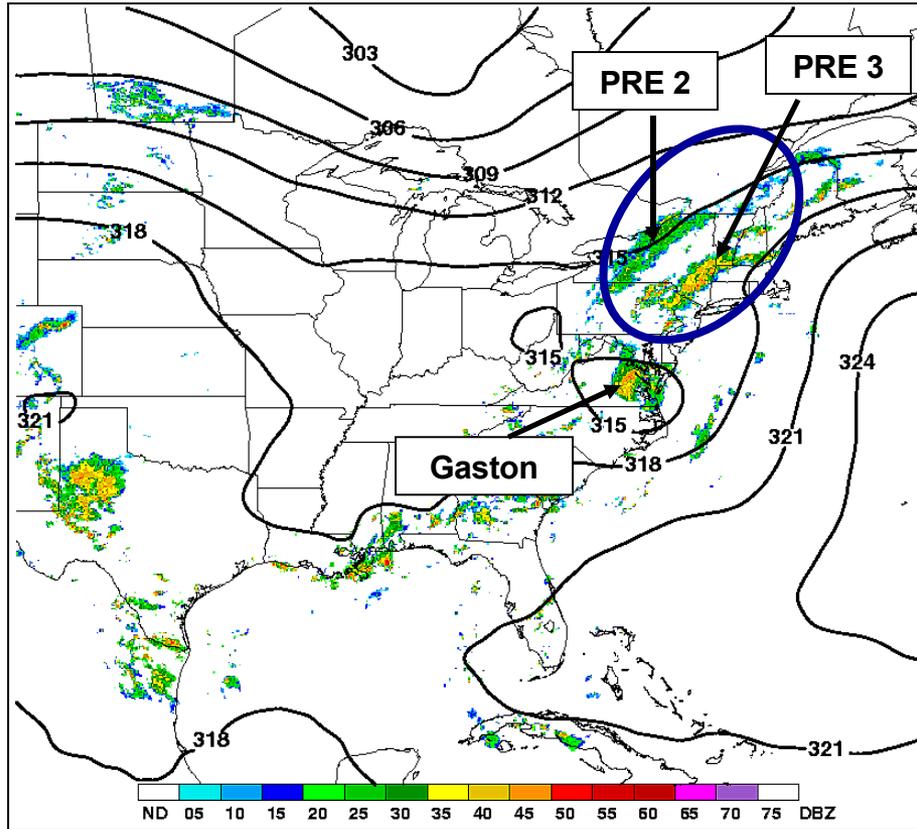


1200 UTC 040830 700 hPa Ht (dam)
WSI NOWRAD image

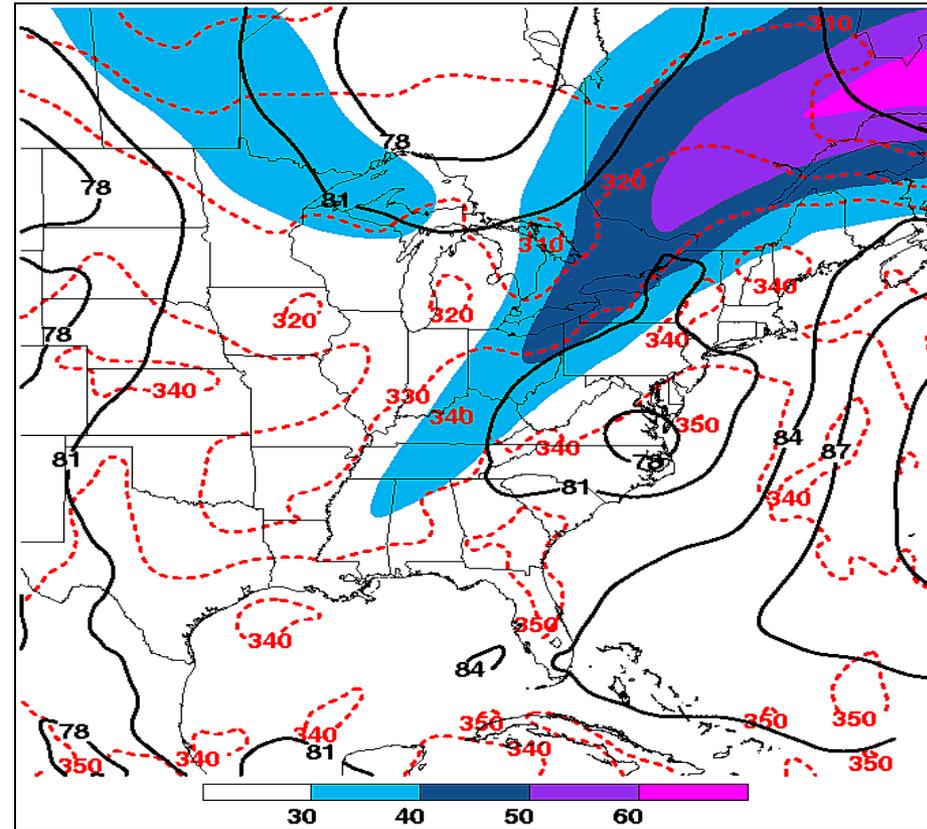


1200 UTC 040830 925 hPa Ht (dam), θ_e (K),
and 200 hPa wind speeds (m s^{-1})

GASTON (2004): SYNOPTIC FEATURES



0000 UTC 040831 700 hPa Ht (dam)
WSI NOWRAD image



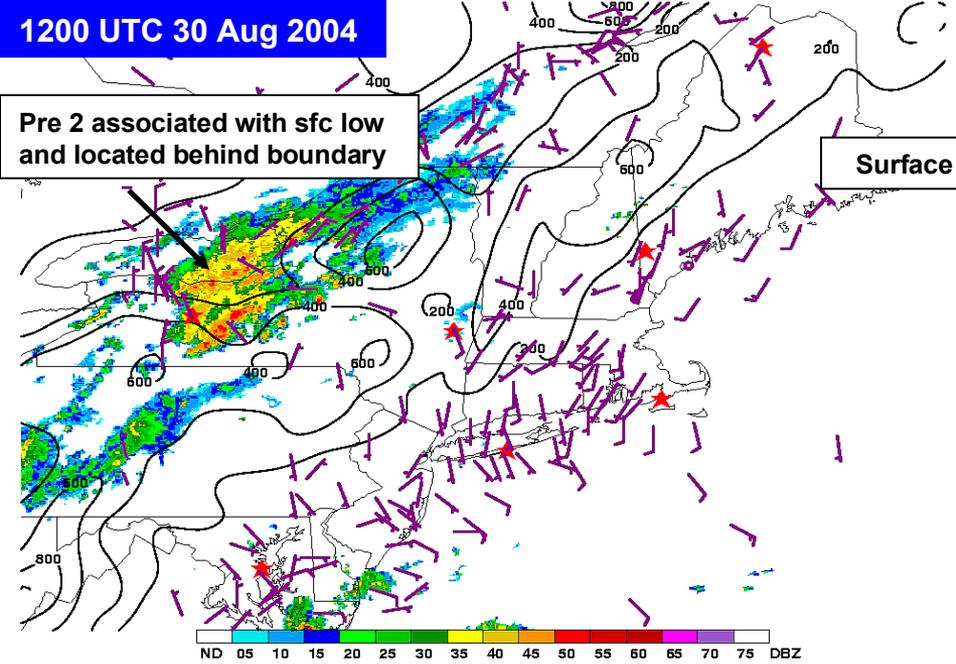
0000 UTC 040831 925 hPa Ht (dam), θ_e (K),
and 200 hPa wind speeds (m s^{-1})

GASTON (2004)

MESOSCALE FEATURES

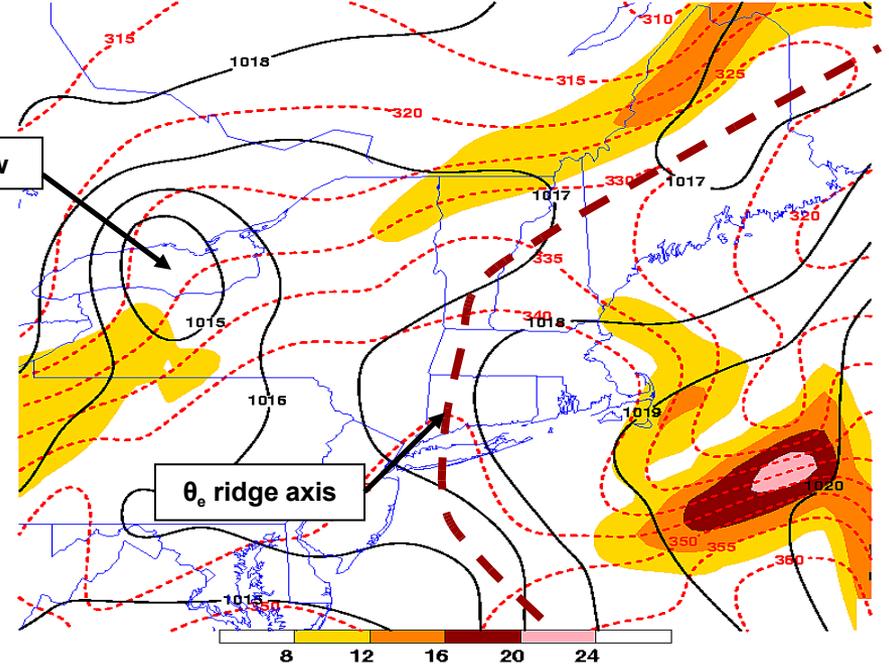
1200 UTC 30 Aug 2004

Pre 2 associated with sfc low and located behind boundary



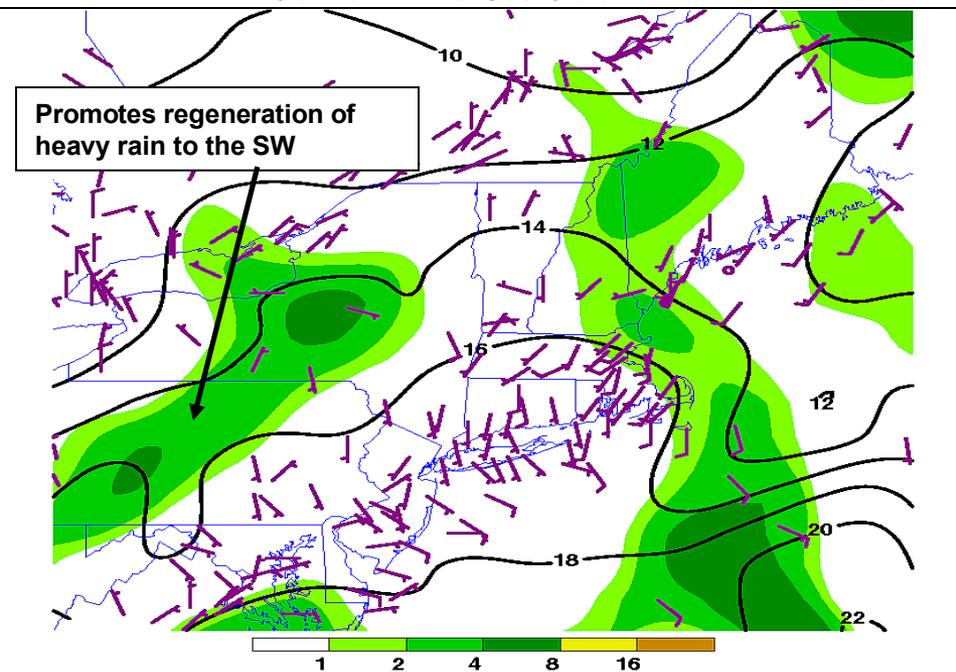
Radar reflectivity (dBZ), RUC topography (m), and sfc wind barbs (kt)

Surface Low



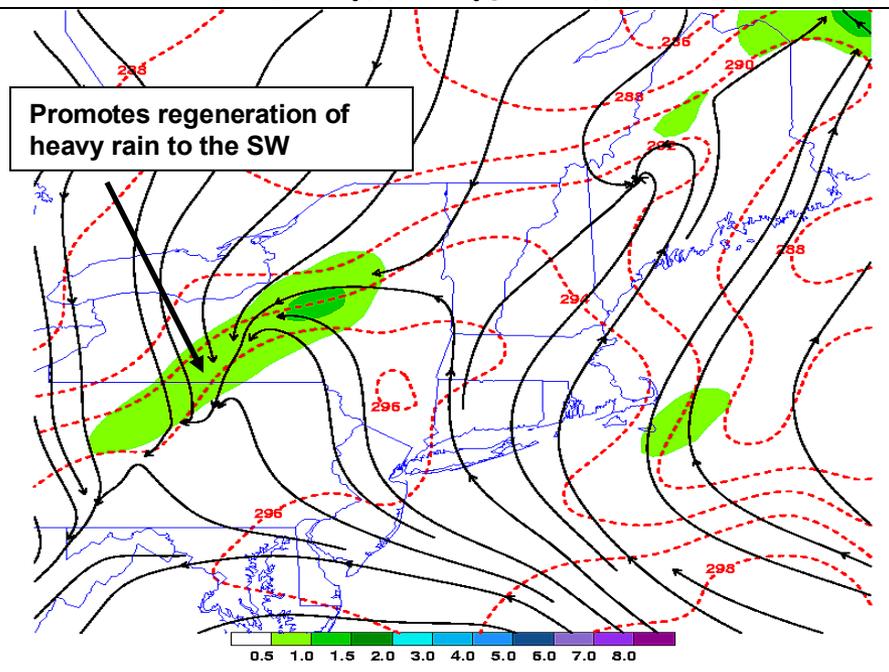
MSLP (hPa), sfc θ_e (K), and θ_e gradient ($K (100 km)^{-1}$)

Promotes regeneration of heavy rain to the SW



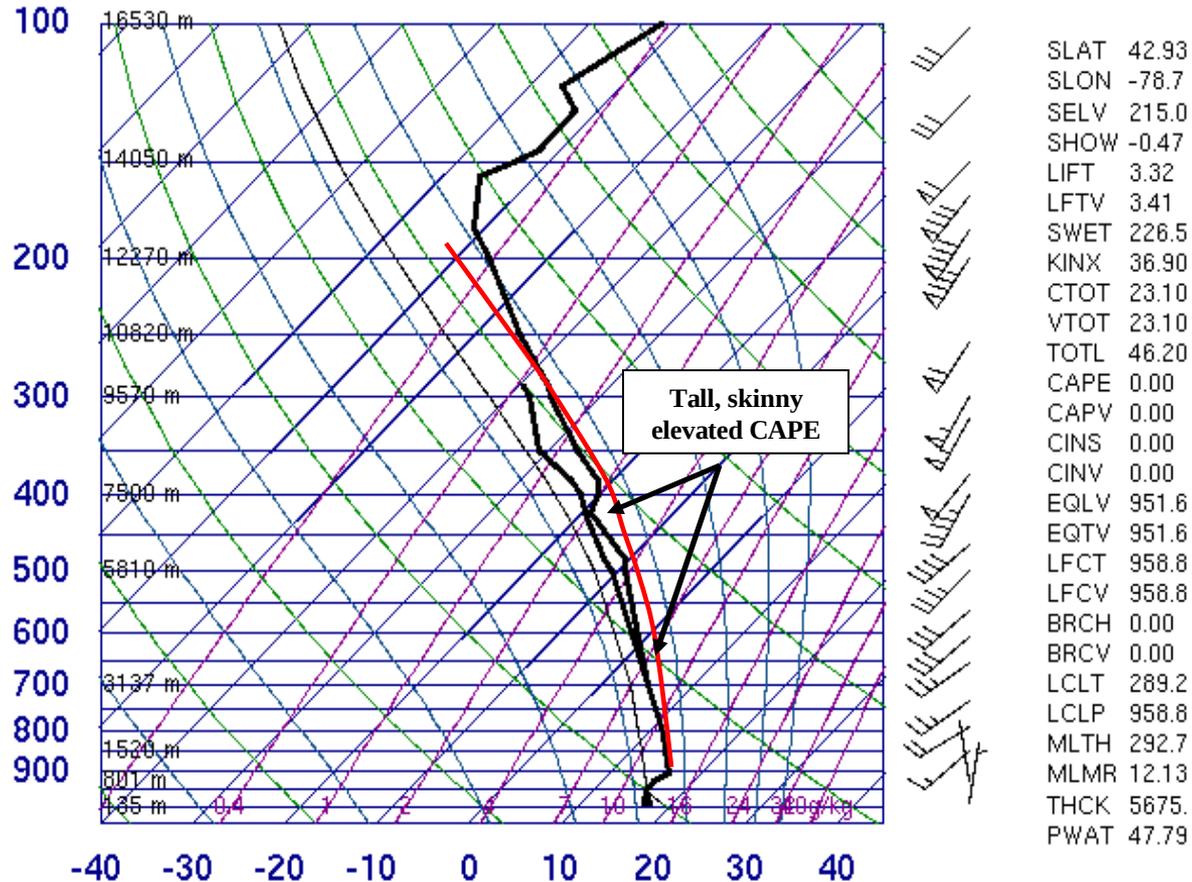
Sfc moisture flux convergence ($10^{-7} s^{-1}$), mixing ratio ($g kg^{-1}$), and wind barbs (kt)

Promotes regeneration of heavy rain to the SW



Sfc frontogenesis ($K (100 km)^{-1} (3 h)^{-1}$), θ_e (K), and streamlines

72528 BUF Buffalo Int

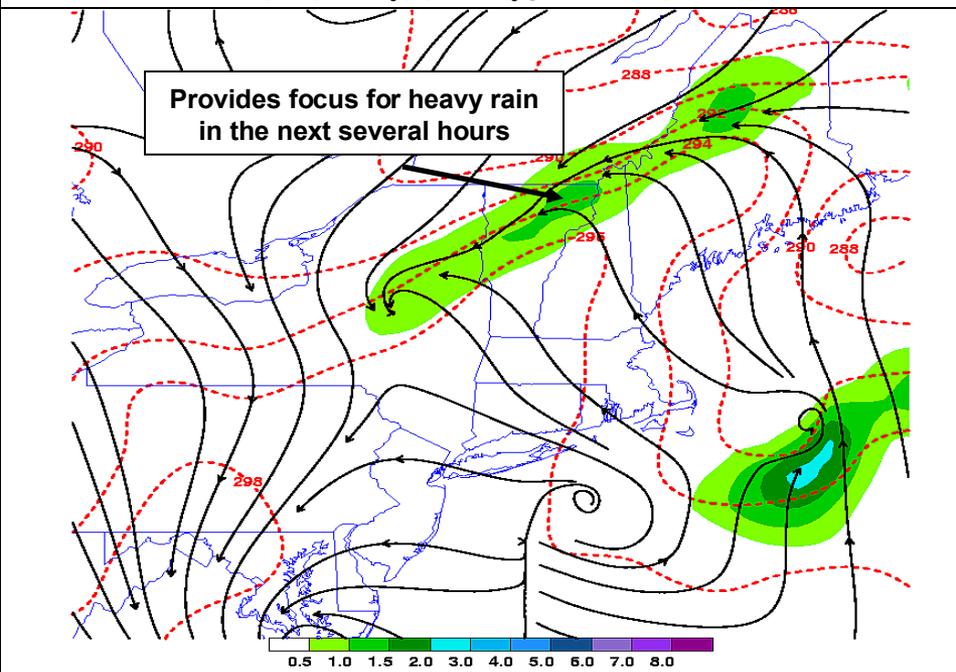
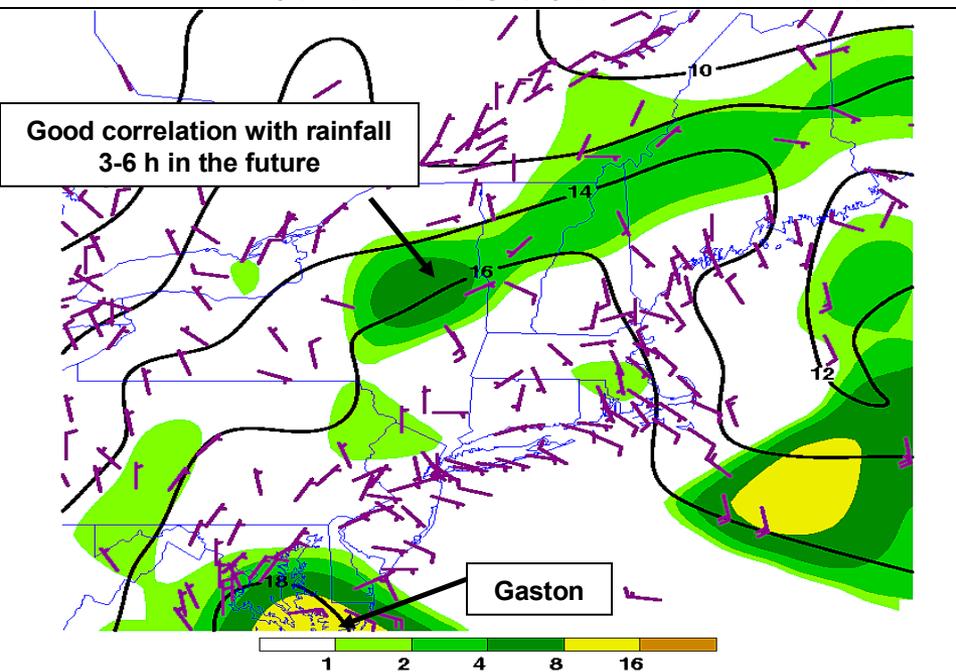
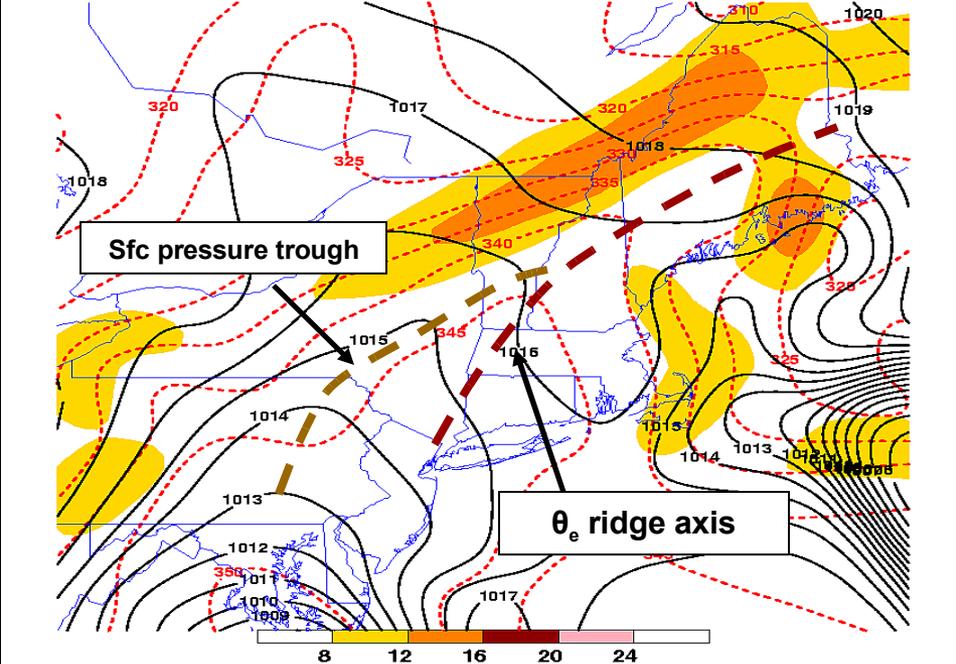
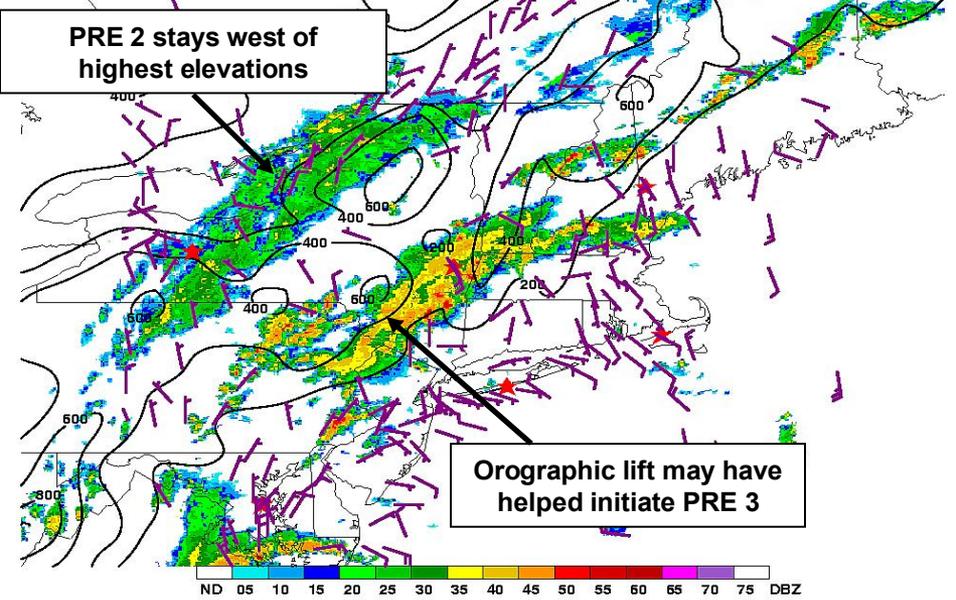


12Z 30 Aug 2004

University of Wyoming

- Saturated up through 600 hPa
- LL northerlies indicate BUF is behind the boundary
- 850-200 hPa speed shear ~65 kt
- K-Index and SWI conducive to heavy convective rainfall

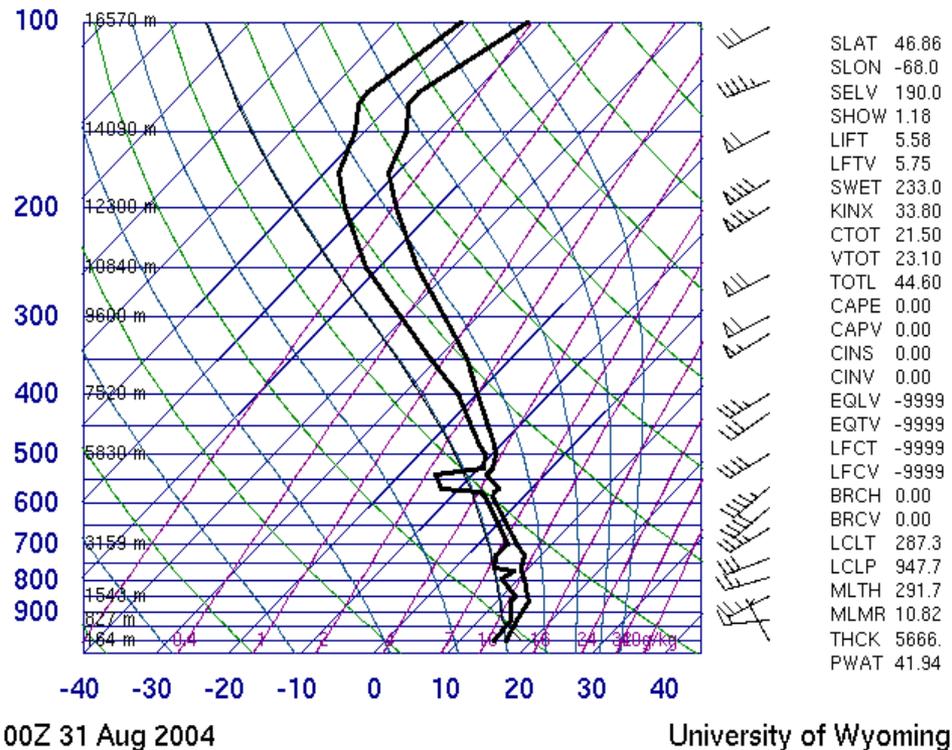
0000 UTC 31 Aug 2004



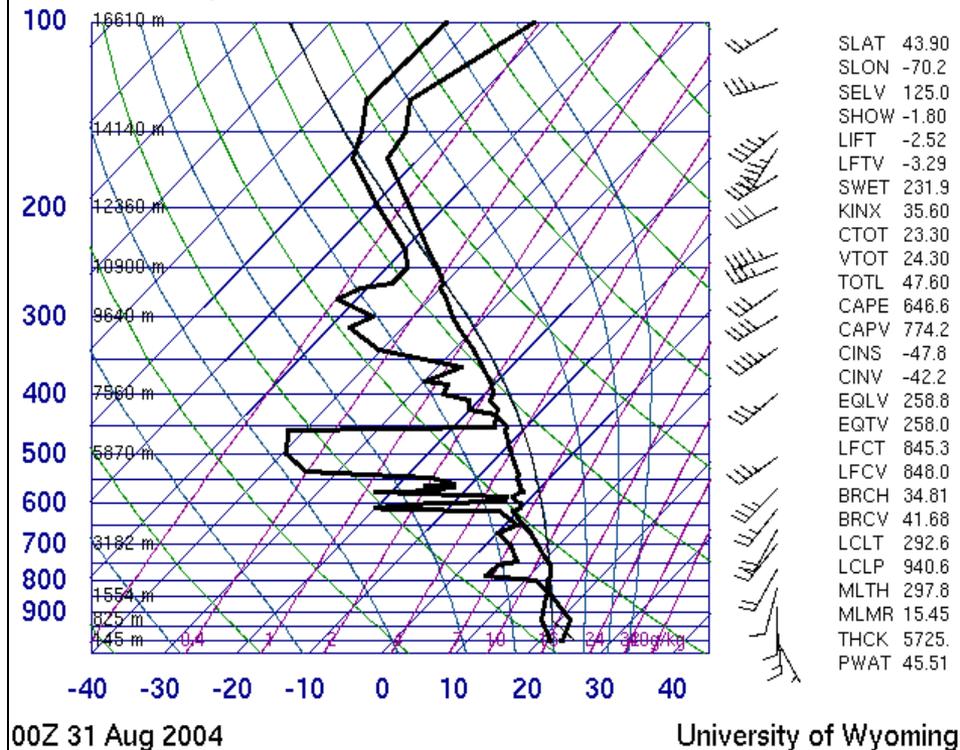
Sfc moisture flux convergence ($10^{-7} s^{-1}$), mixing ratio ($g kg^{-1}$), and wind barbs (kt)

Sfc frontogenesis ($K (100 km)^{-1} (3 h)^{-1}$), θ (K), and streamlines

72712 CAR Caribou



74389 GYX Gray



- PRE 3 affects region several hours later
- Located behind θ_e boundary
- 850-200 hPa speed shear ~50 kt
- K-Index still conducive to heavy convective rainfall

- Far eastern edge of PRE 3 affects region ~2 h later
- Located ahead of θ_e boundary
- LL veering winds with surface-based CAPE

CONCLUSIONS

- **~ 1/3 of all US landfalling TCs produce at least one PRE, but landfall is not necessary**
- **PREs form on the order of 1000 km away from their parent TCs and about 1-2 days in advance**
- **LOT PREs are most common, but AT and ROT PREs produce the highest rain rates**
- **TCs recurving over the Southeast and along the East Coast have the greatest likelihood of producing PREs**

CONCLUSIONS

PREs generally form:

- When persistent, deep meridional flow transports tropical air far from the TC**
- In favored upslope regions or along synoptic/mesoscale boundaries:**
 - Ahead of the boundary, where surface-based convection is favored**
 - Immediately behind the boundary, where elevated convection is favored**

CONCLUSIONS

PREs generally form:

- Along and just west of a low-level θ_e ridge, near the strongest gradient
- Near a midlevel jet-entrance region confluence zone
- Under favorable upper-level jet dynamics

TECHNOLOGY TRANSFER

GETTING THE WORD OUT

- Research presented at BGM Spring Workshop
27 March 2007
- Other conference presentations
- Possible NWS teletraining
- Eventual journal publication

TECHNOLOGY TRANSFER

INCORPORATING INTO OPERATIONS

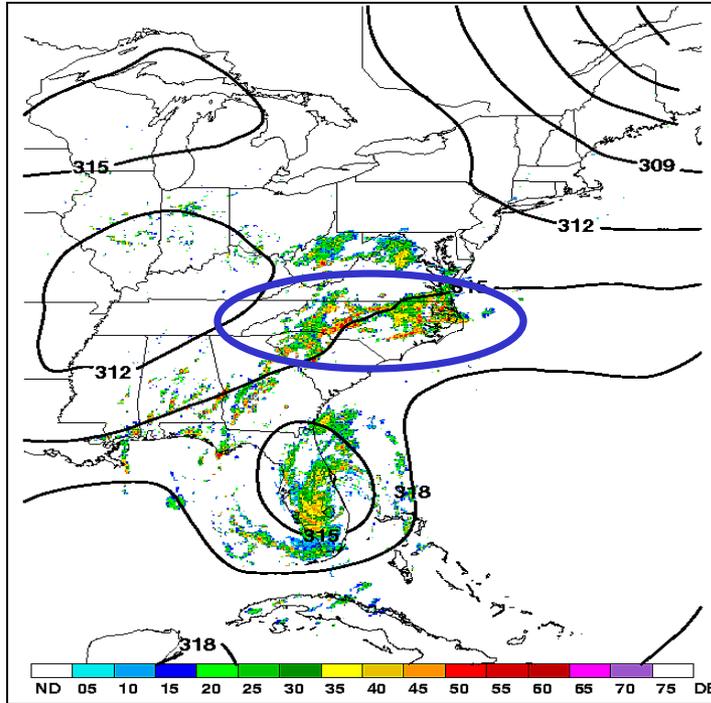
- **24-h track/precip technique for capturing rainfall downstream of TC**
- **TC track climatological maps**
- **Statistically modified Bosart and Carr (1978) conceptual model**
- **Ingredients-based methodology**
- **Construction of an “all-in-one” conceptual model**

QUESTIONS?

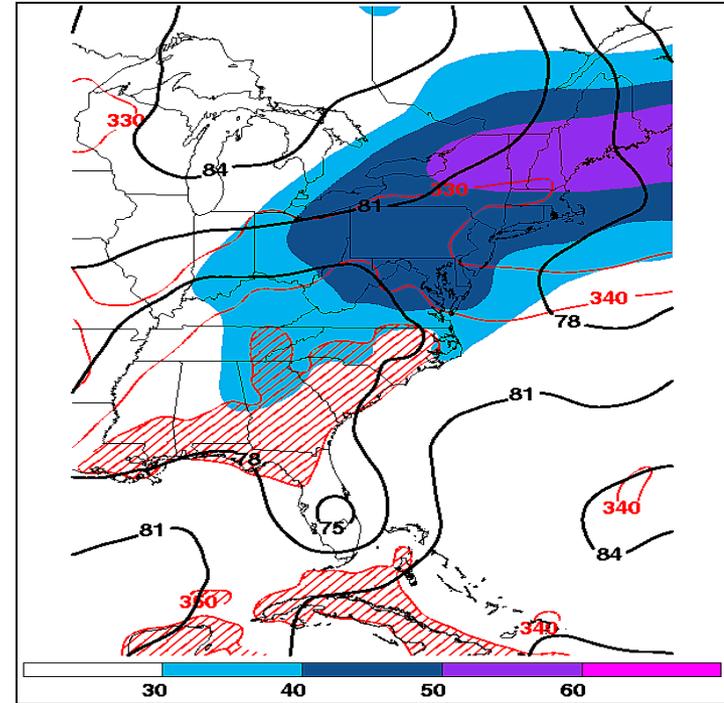
COMMENTS?

mcote@atmos.albany.edu

ALONG TRACK PREs



2100 UTC 060830 700 hPa Ht (dam)
WSI NOWRAD image

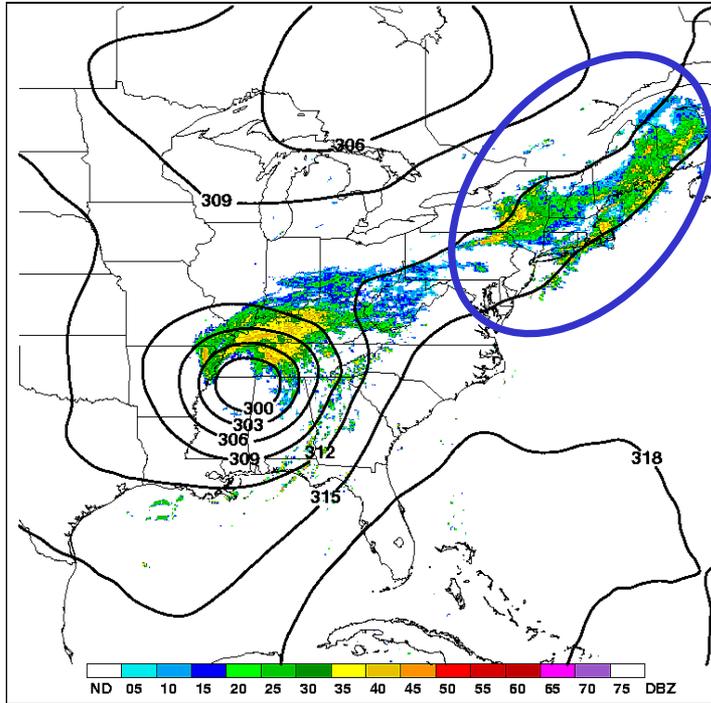


2100 UTC 060830 925 hPa Ht (dam), θ_e (K),
200 hPa wind speed (m s^{-1}) and

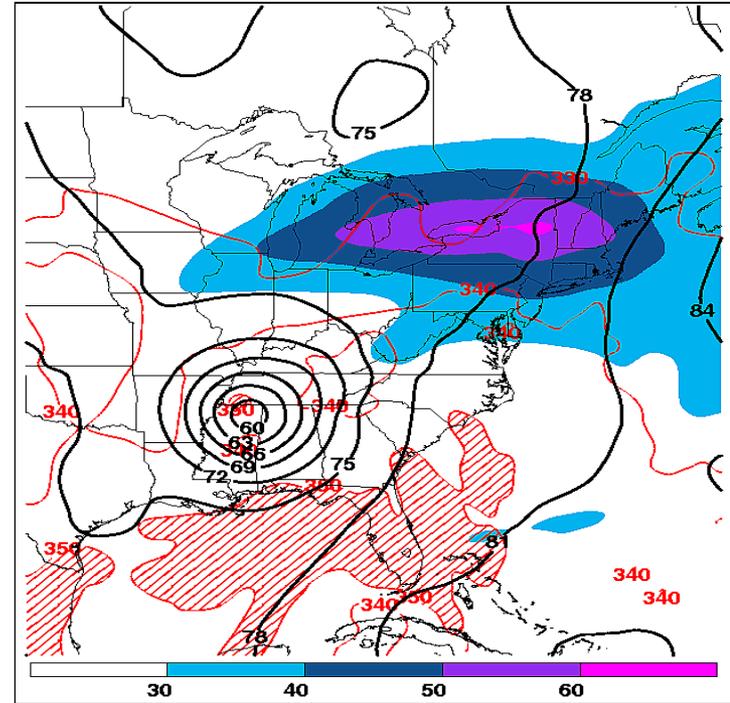
Ernesto (2006)

- NW/SE oriented trough well to the northeast
- Closed midlevel low NW and flat ridge east of TC
- Broad upper-level jet to the north
- On western edge of θ_e ridge

RIGHT OF TRACK PREs



0900 UTC 050830 700 hPa Ht (dam)
WSI NOWRAD image

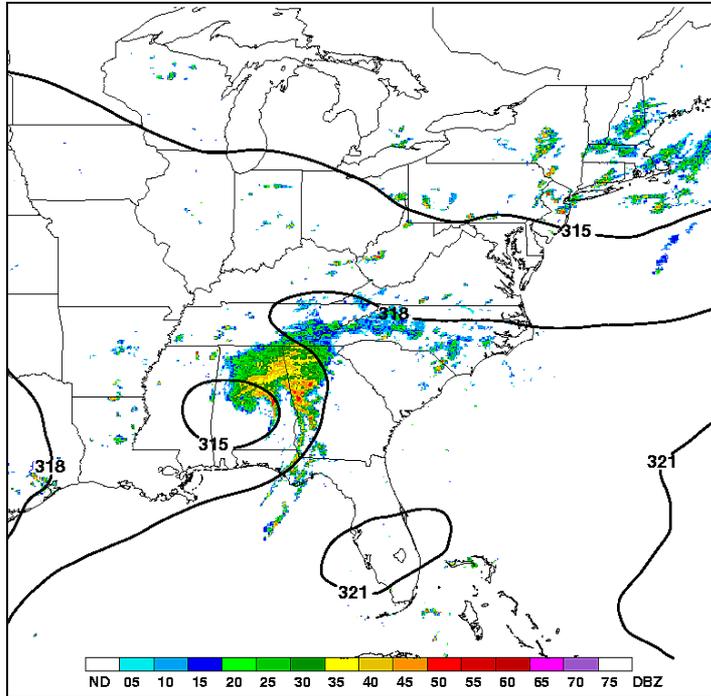


0900 UTC 050830 925 hPa Ht (dam), θ_e (K),
200 hPa wind speed (m s⁻¹) and

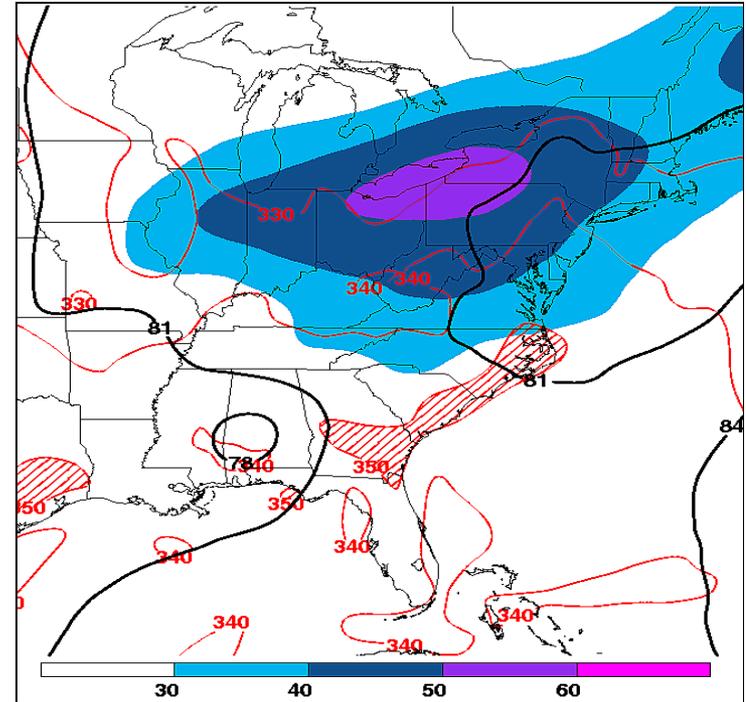
Katrina (2005)

- Large midlevel low NNE and ridge SE of TC
- PREs a bit downstream of where model predicts
- Jet dynamics only partially explain the PREs
- No prominent low-level θ_e ridge or gradient near PRE

NULL CASE



0000 UTC 050707 700 hPa Ht (dam)
WSI NOWRAD image



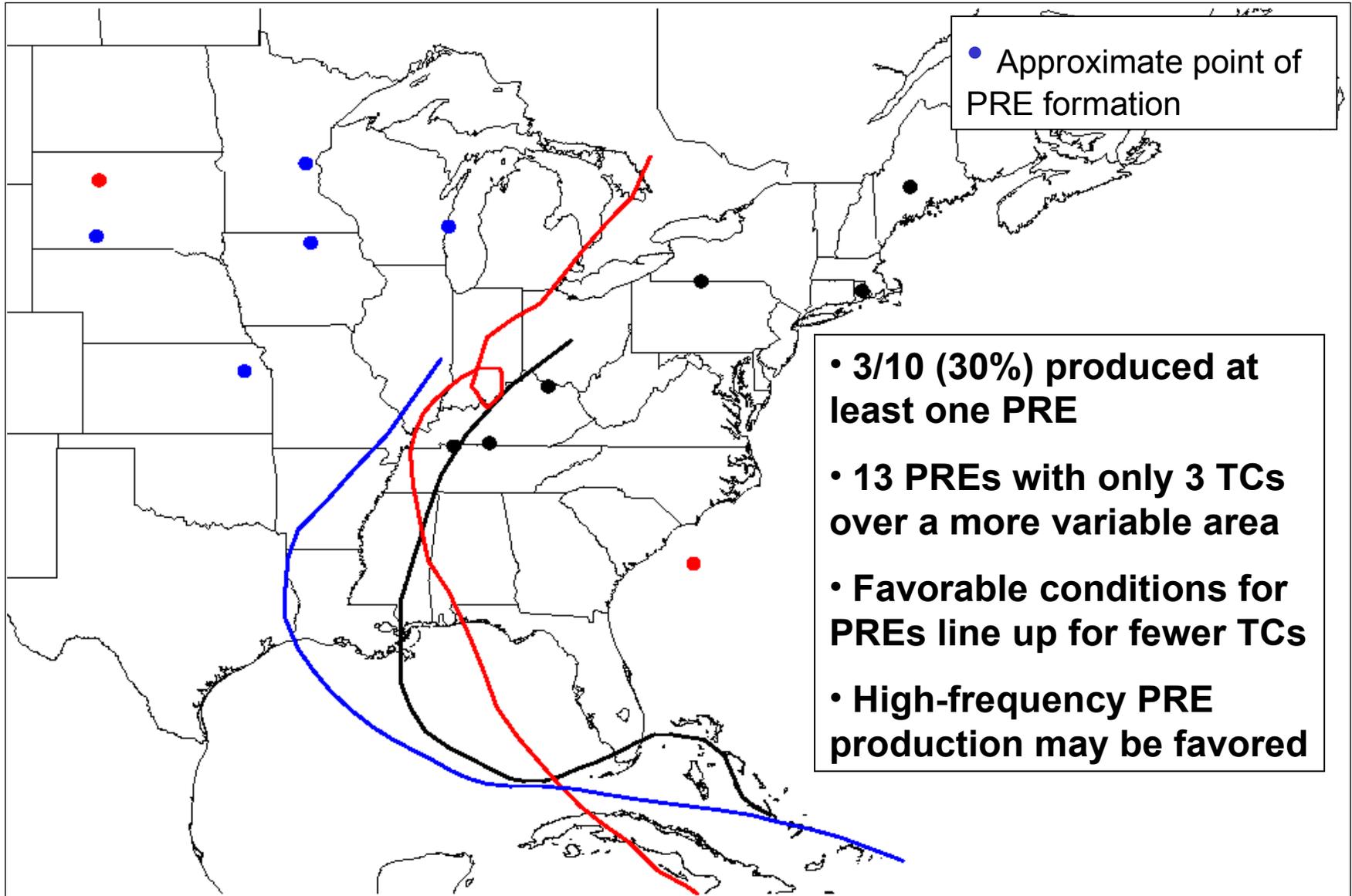
0000 UTC 050707 925 hPa Ht (dam), θ_e (K),
200 hPa wind speed (m s⁻¹)

and

Cindy (2005)

- WNW flow at midlevels
- Scattered rainfall over New England not related to Cindy
- Massive low-level ridge poleward of TC
- No rainfall near low-level θ_e ridge

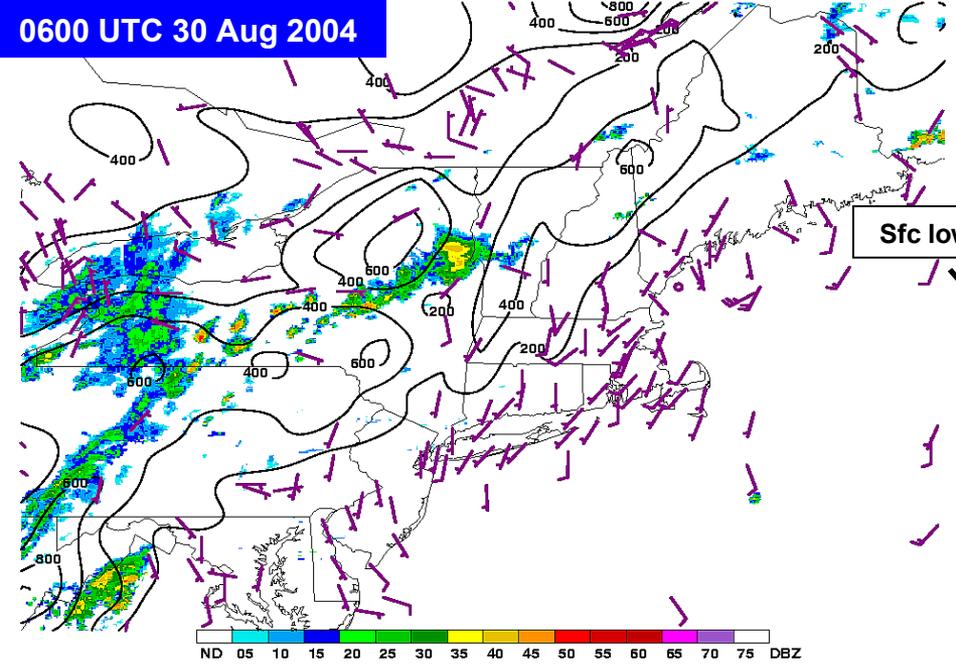
CENTRAL GULF LANDFALLS



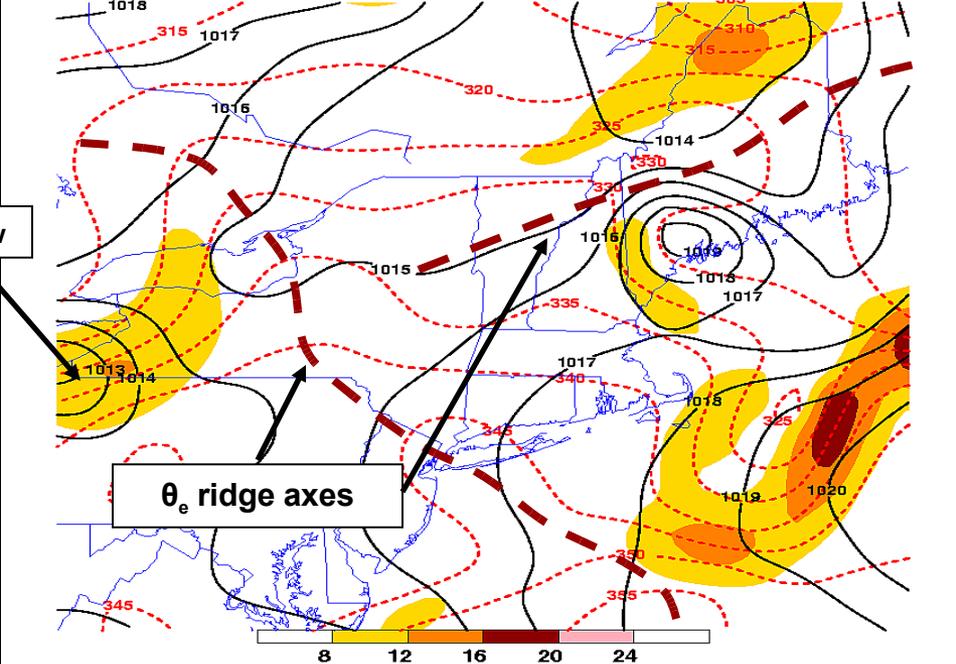
GASTON (2004)

**MESOSCALE DETAILS FROM
ADDITIONAL TIMES**

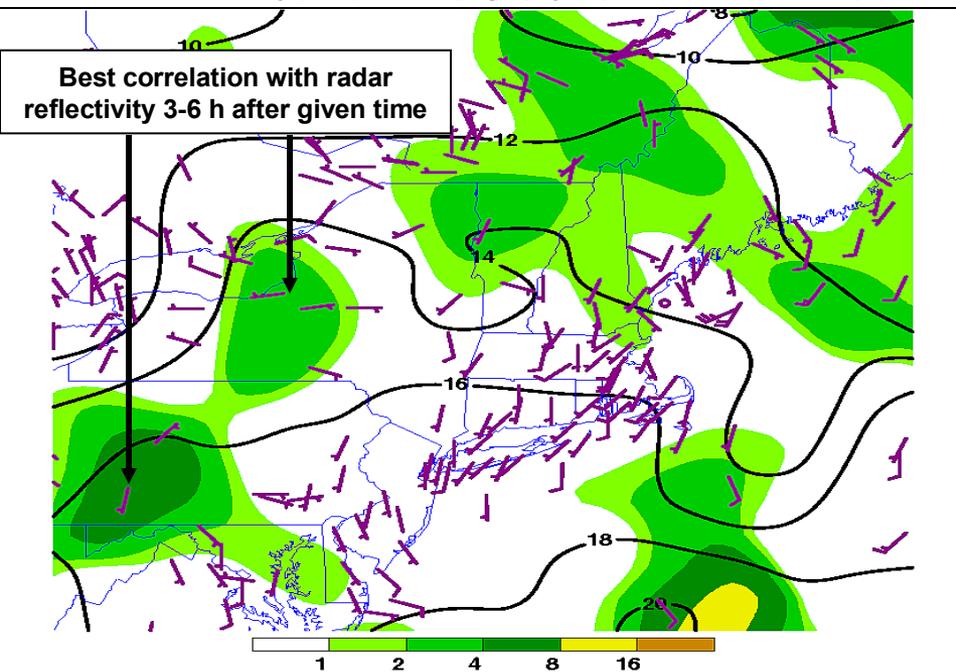
0600 UTC 30 Aug 2004



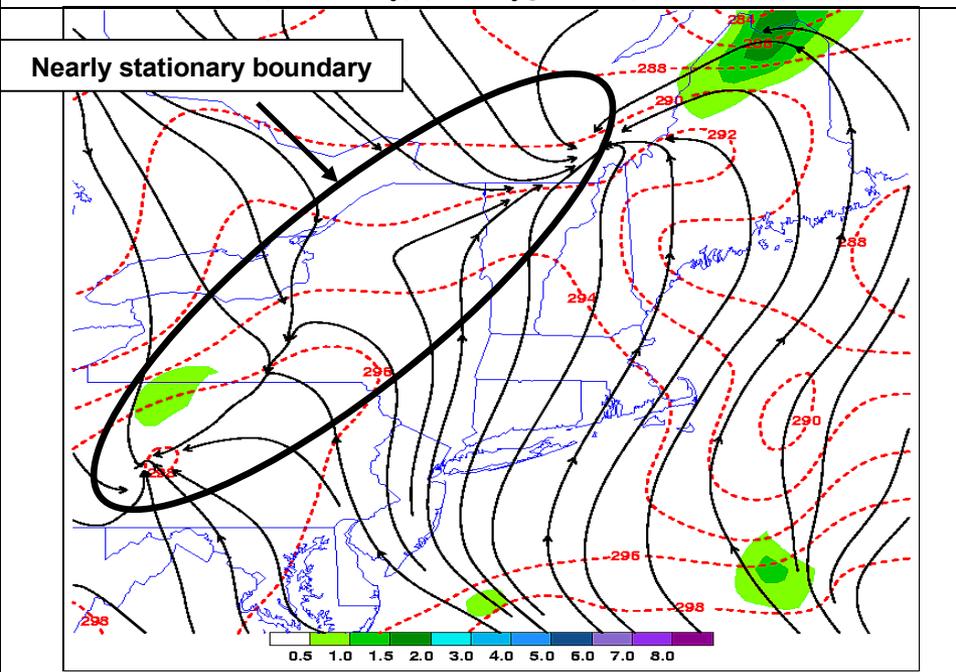
Radar reflectivity (dBZ), RUC topography (m), and sfc wind barbs (kt)



MSLP (hPa), sfc θ_e (K), and θ_e gradient ($K (100 km)^{-1}$)



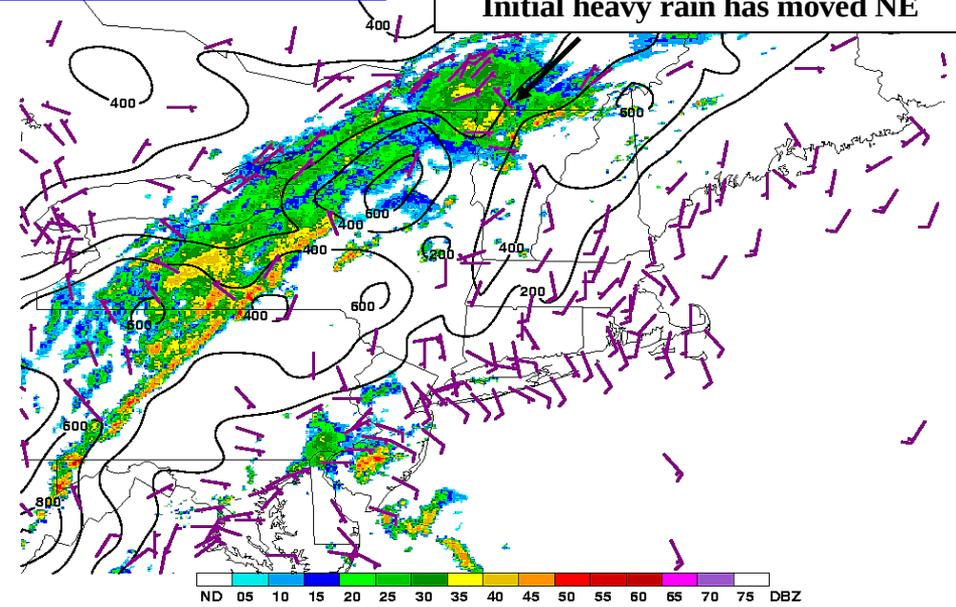
Sfc moisture flux convergence ($10^{-7} s^{-1}$), mixing ratio ($g kg^{-1}$), and wind barbs (kt)



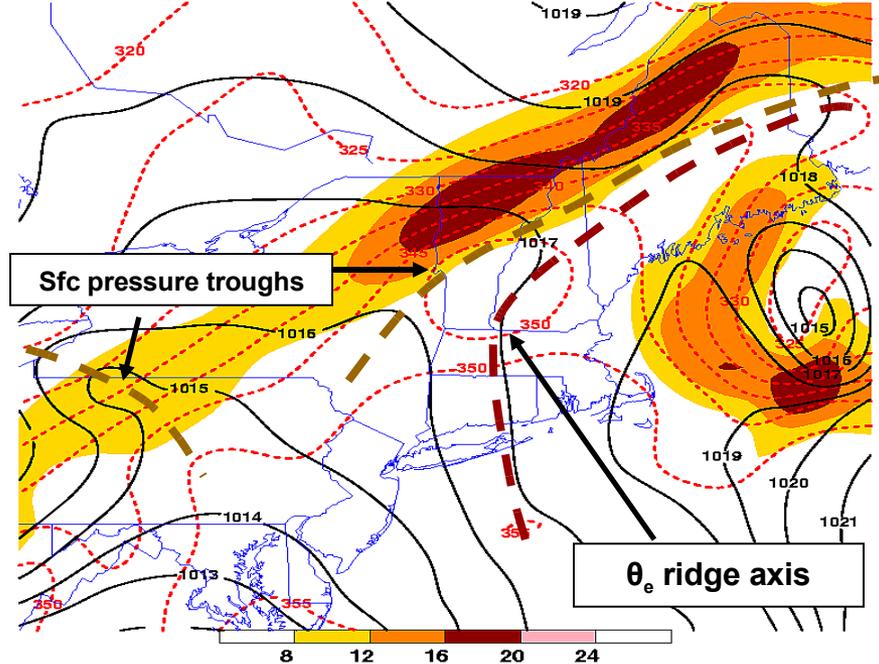
Sfc frontogenesis ($K (100 km)^{-1} (3 h)^{-1}$), θ (K), and streamlines

1800 UTC 30 Aug 2004

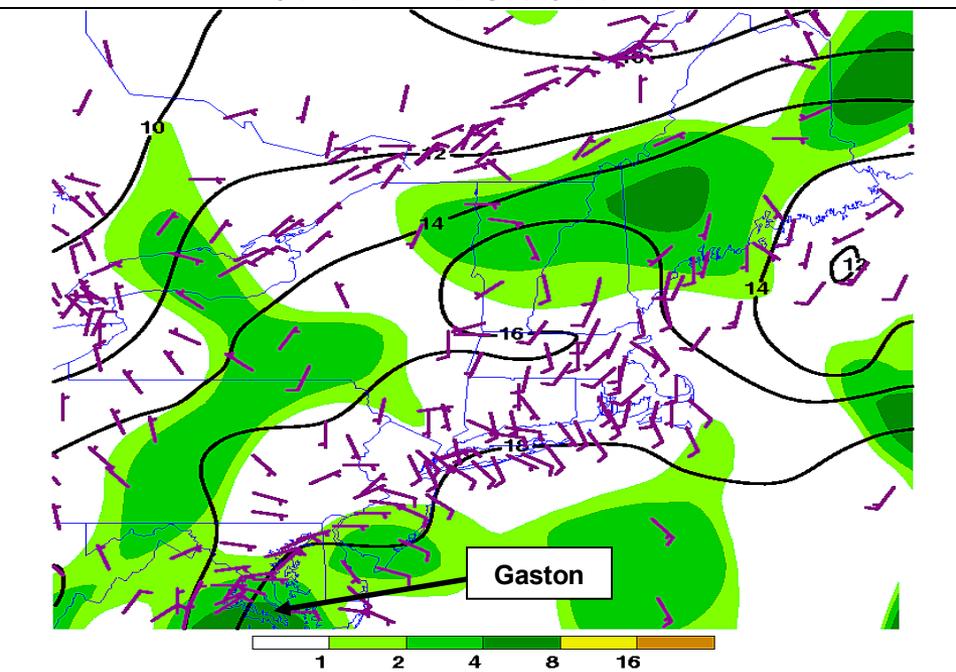
Initial heavy rain has moved NE



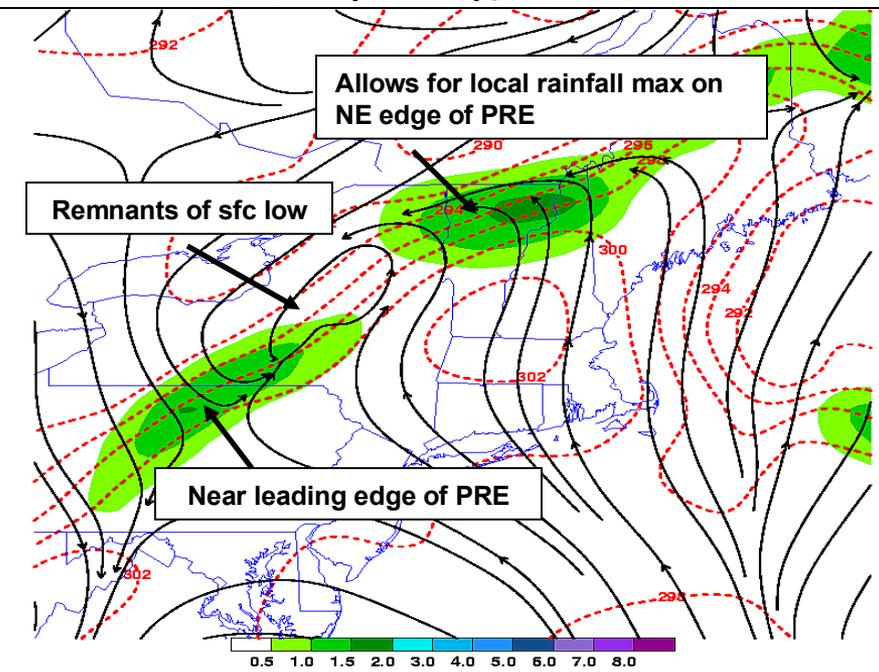
Radar reflectivity (dBZ), RUC topography (m), and sfc wind barbs (kt)



MSLP (hPa), sfc θ_e (K), and θ_e gradient ($K (100 km)^{-1}$)

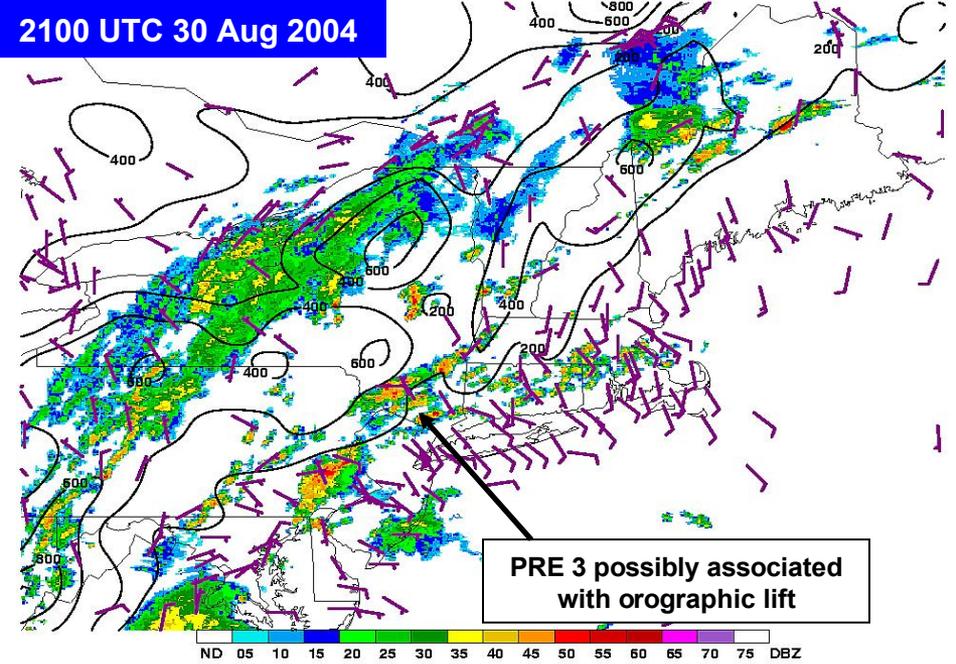


Sfc moisture flux convergence ($10^{-7} s^{-1}$), mixing ratio ($g kg^{-1}$), and wind barbs (kt)

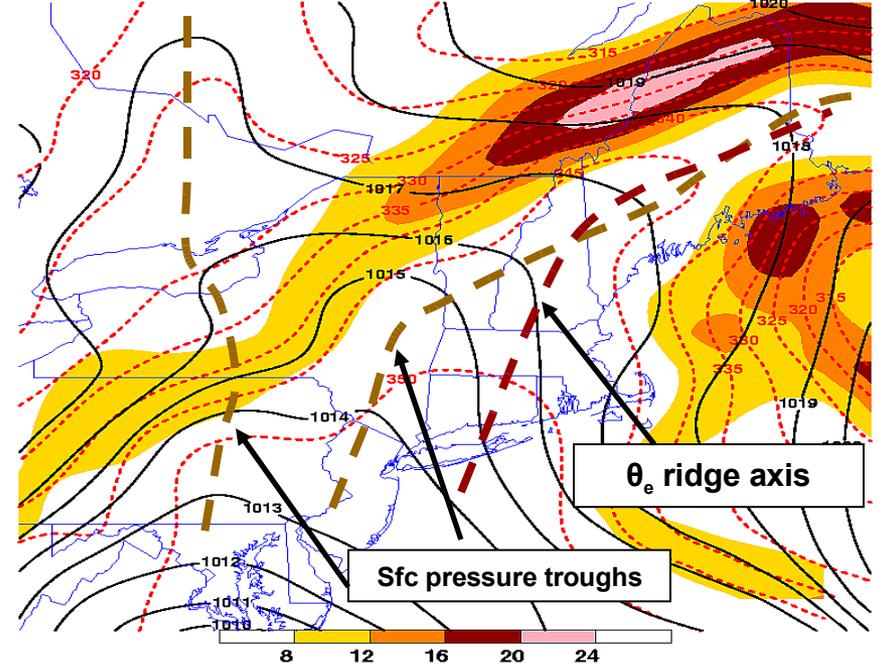


Sfc frontogenesis ($K (100 km)^{-1} (3 h)^{-1}$), θ (K), and streamlines

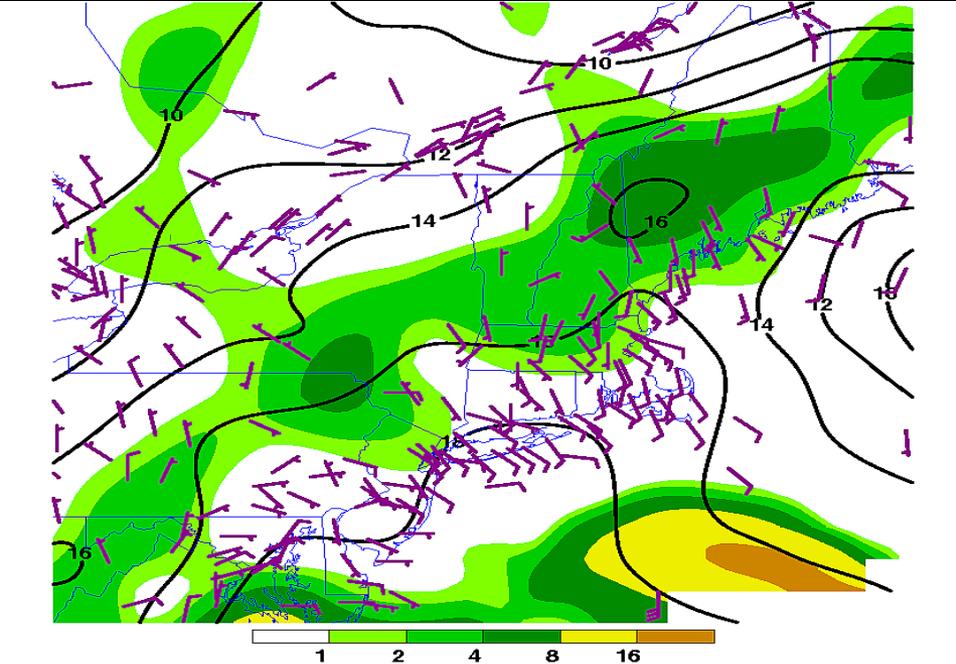
2100 UTC 30 Aug 2004



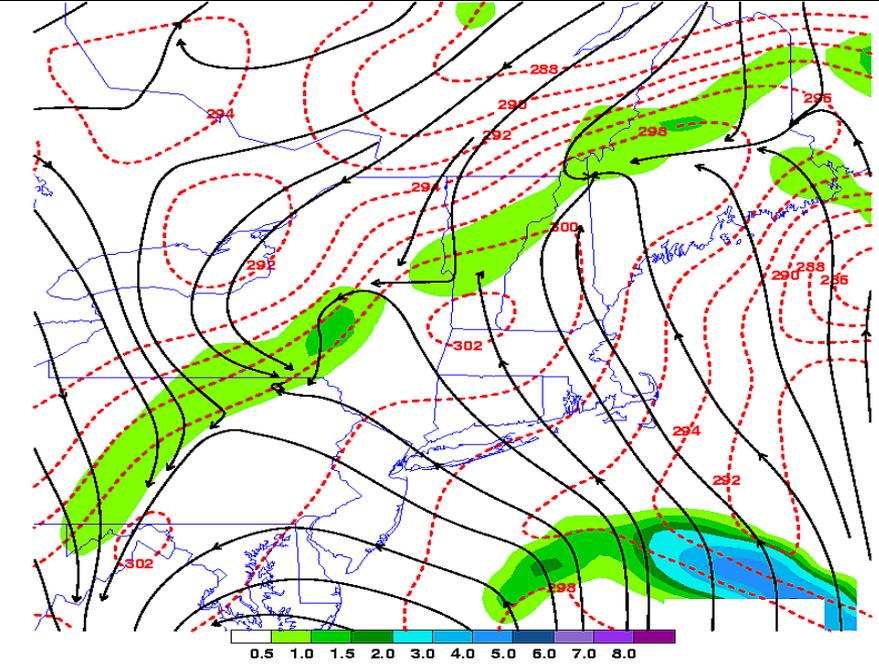
Radar reflectivity (dBZ), RUC topography (m), and sfc wind barbs (kt)



MSLP (hPa), sfc θ_e (K), and θ_e gradient ($K (100 km)^{-1}$)

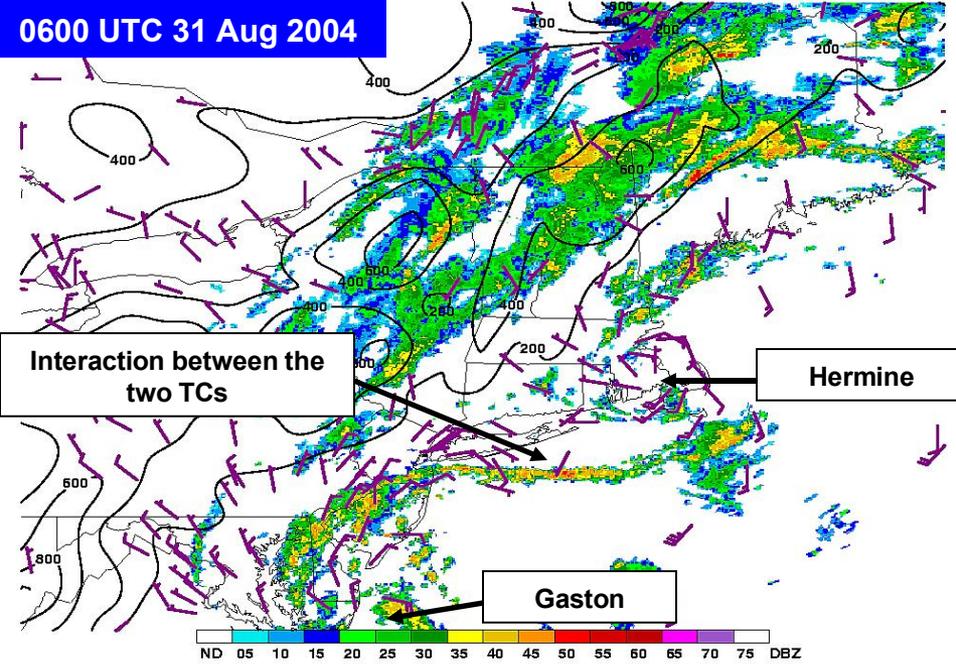


Sfc moisture flux convergence ($10^{-7} s^{-1}$), mixing ratio ($g kg^{-1}$), and wind barbs (kt)

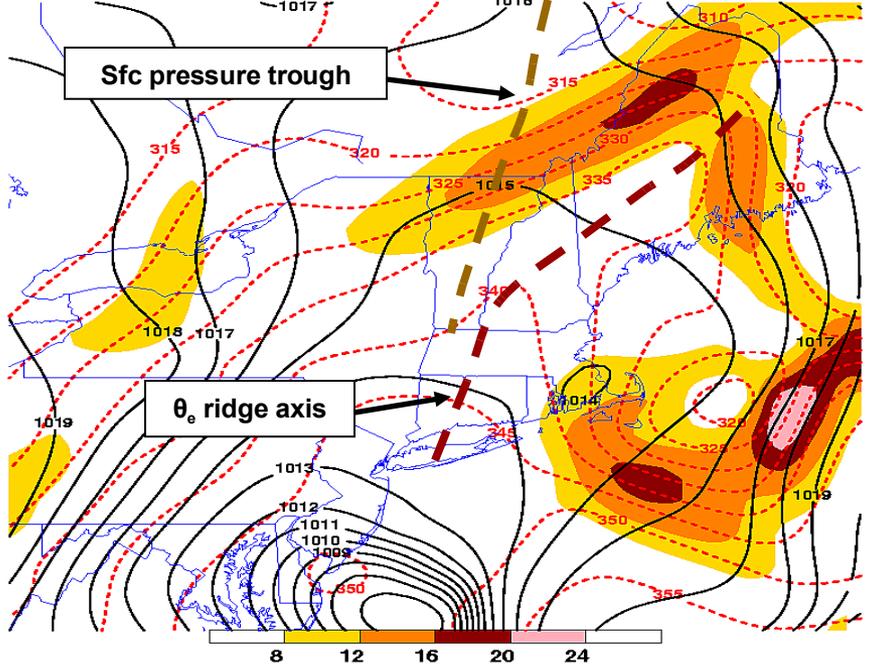


Sfc frontogenesis ($K (100 km)^{-1} (3 h)^{-1}$), θ (K), and streamlines

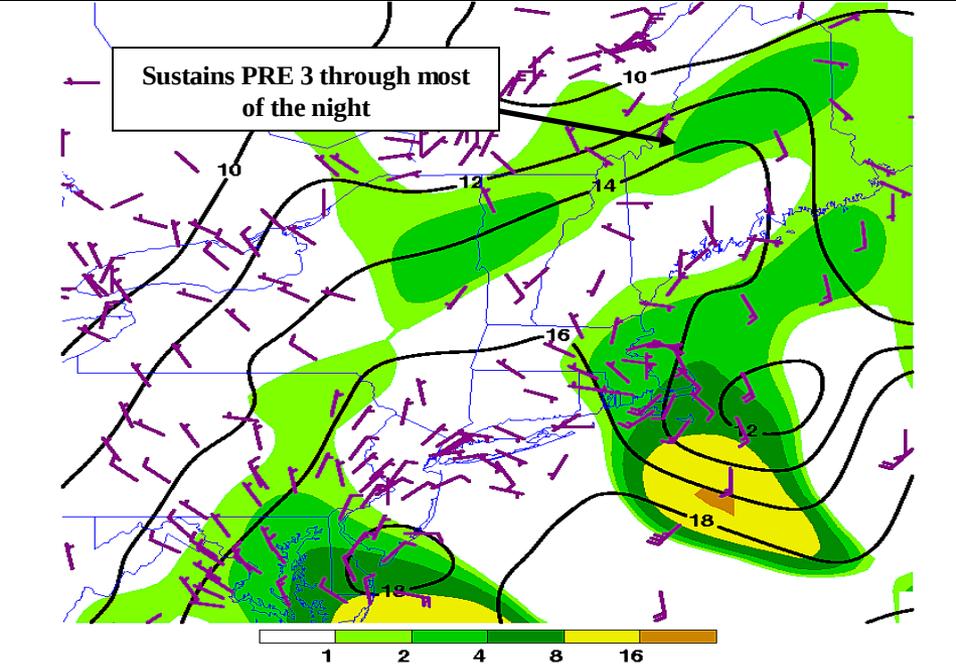
0600 UTC 31 Aug 2004



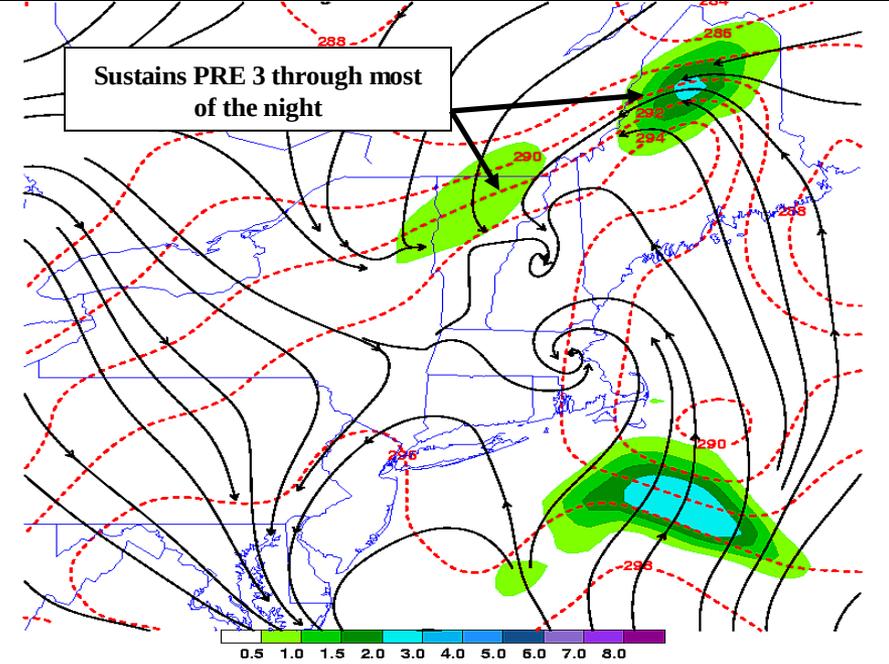
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MSLP (hPa), sfc θ_e (K), and θ_e gradient ($\text{K} (100 \text{ km})^{-1}$)

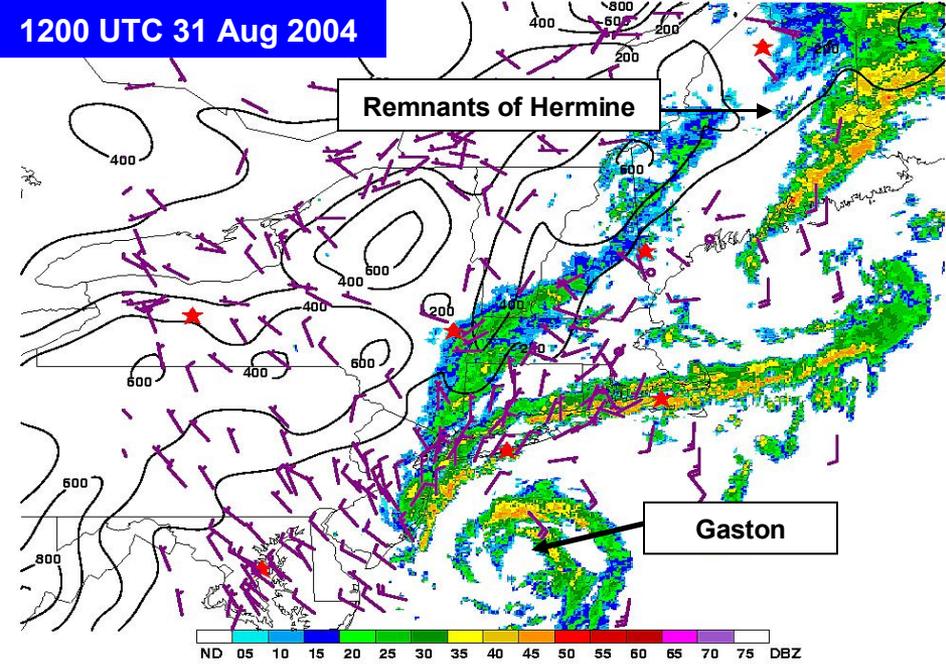


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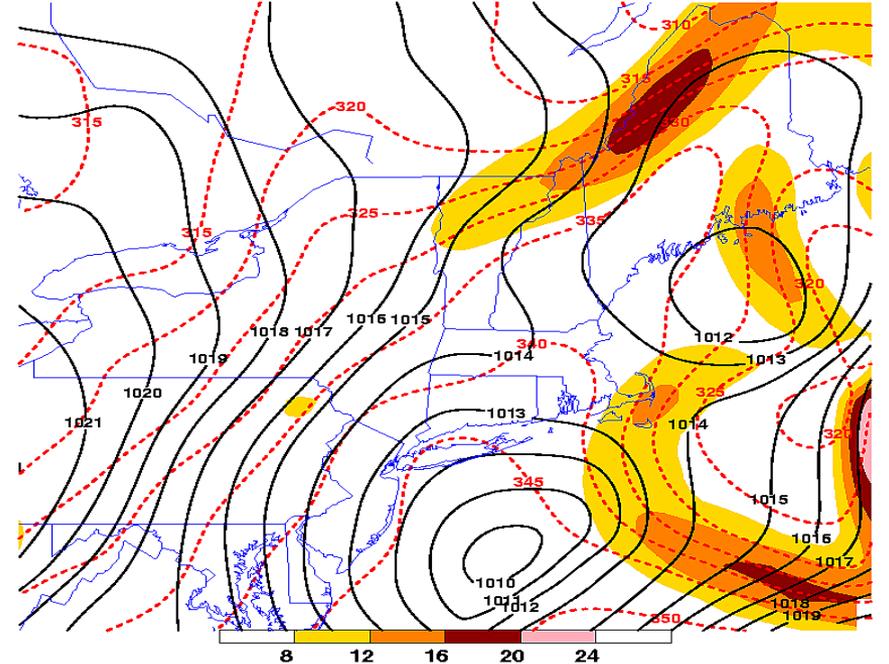


Sfc frontogenesis ($\text{K} (100 \text{ km})^{-1} (3 \text{ h})^{-1}$), θ (K), and streamlines

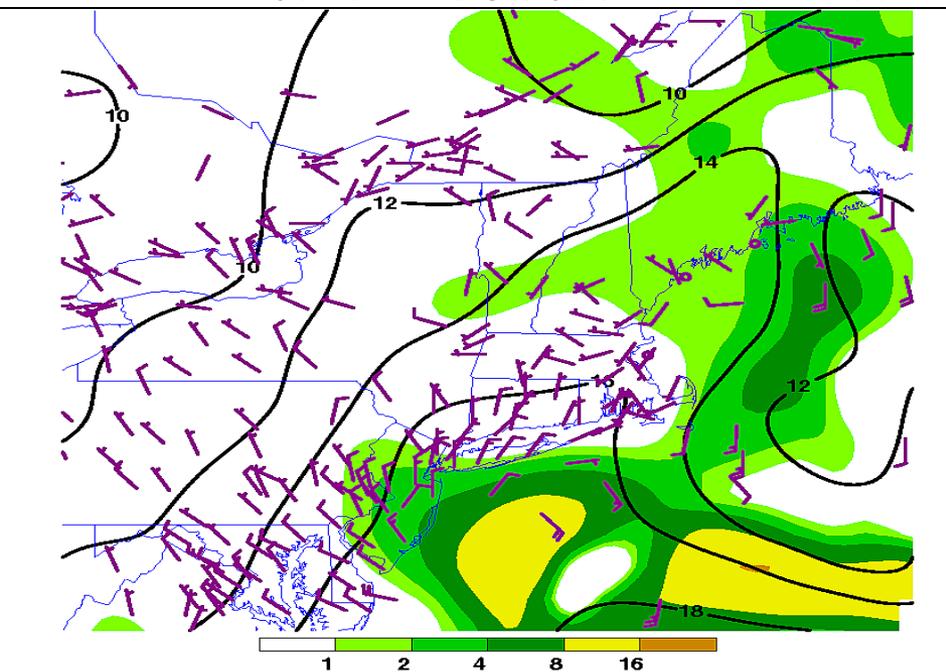
1200 UTC 31 Aug 2004



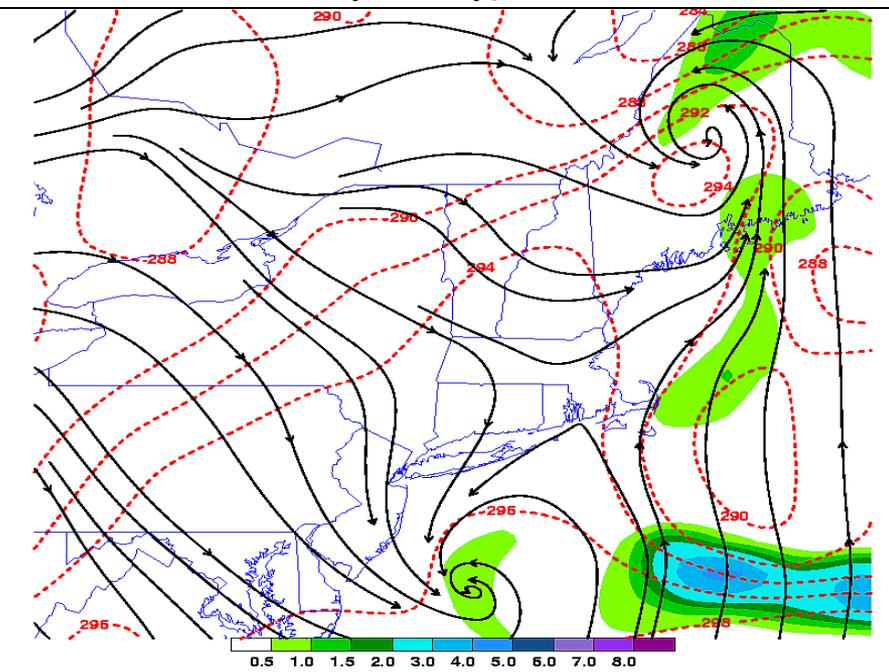
Radar reflectivity (dBZ), RUC topography (m), and sfc wind barbs (kt)



MSLP (hPa), sfc θ_s (K), and θ_s gradient ($K (100 km)^{-1}$)



Sfc moisture flux convergence ($10^{-7} s^{-1}$), mixing ratio ($g kg^{-1}$), and wind barbs (kt)



Sfc frontogenesis ($K (100 km)^{-1} (3 h)^{-1}$), θ (K), and streamlines